

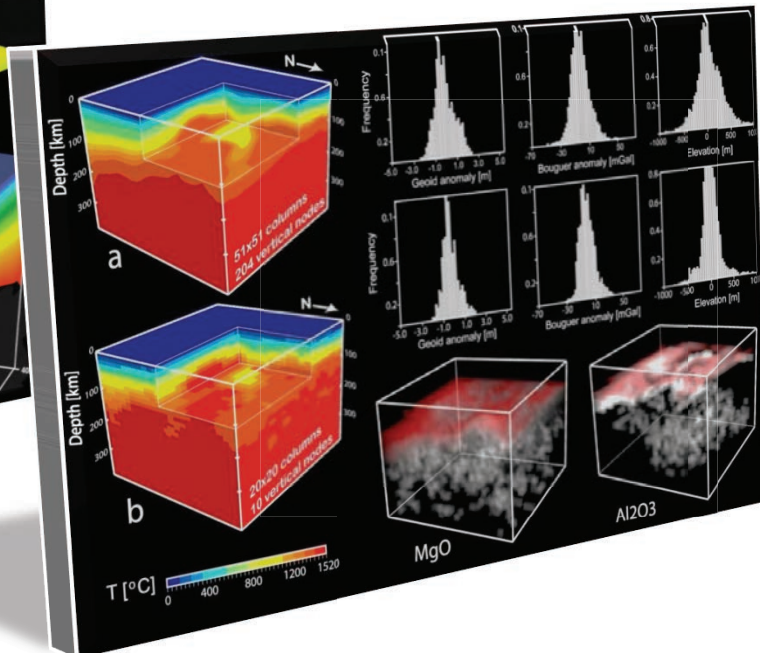
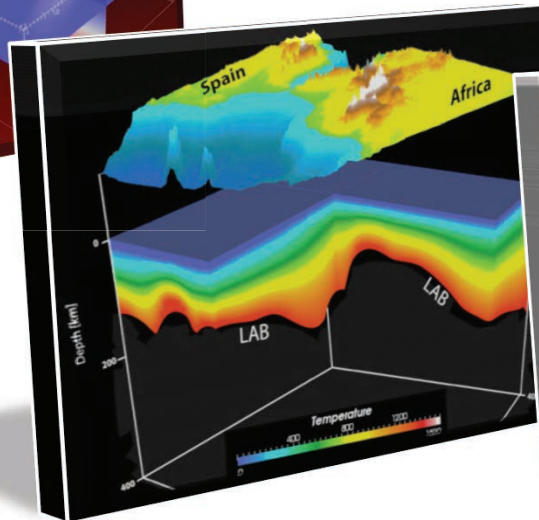
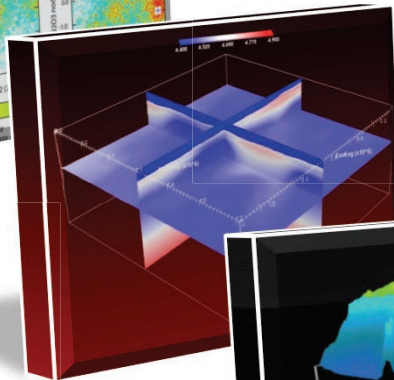
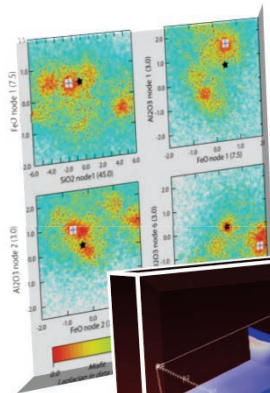
Thermochemical Tomography of the Lithosphere from Multi-observable Probabilistic Inversions

Juan Carlos Afonso

CCFS-GEMOC, Macquarie University, Sydney, Australia.

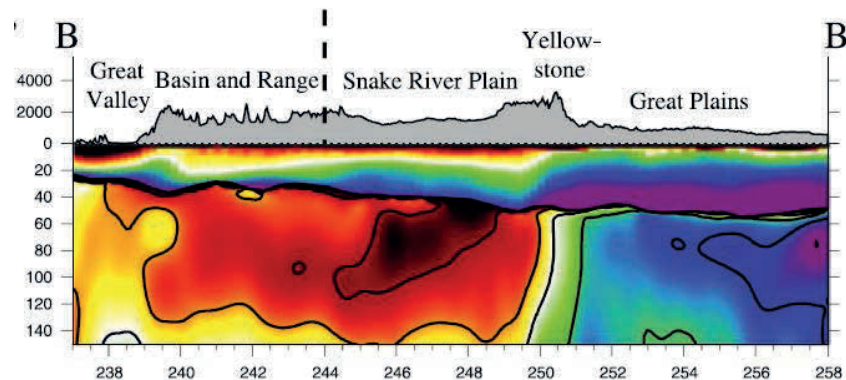
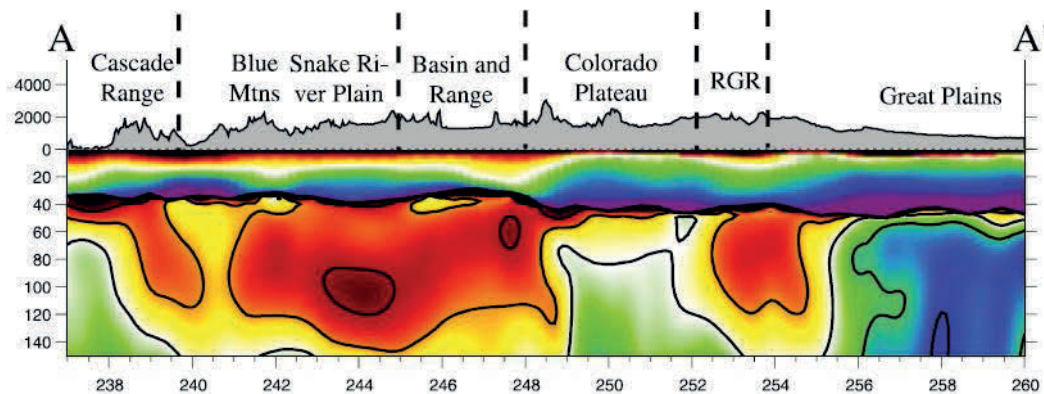
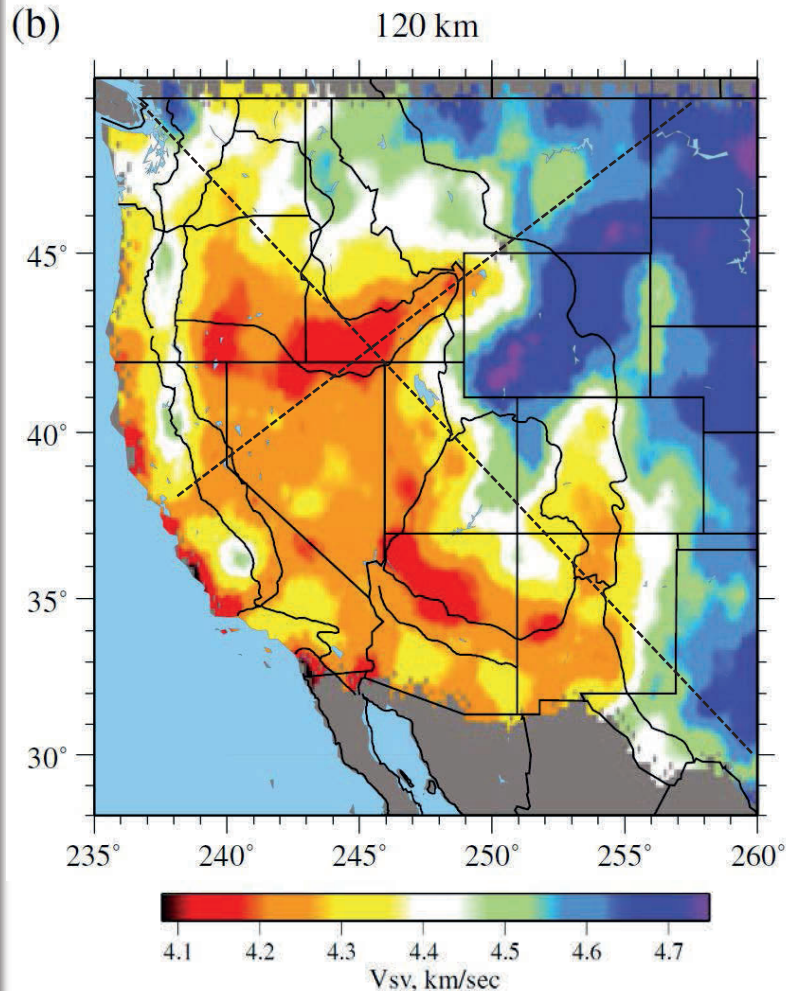
In collaboration with:

Javier Fullea, Yingjie Yang, Nick Rawlinson,
James Connolly, Alan G. Jones, Bill Griffin and
Sue O'Reilly



1) An interesting problem

- What is the nature of the heterogeneity “observed” in the mantle?





1) *An interesting problem*

- What is the nature of the heterogeneity “observed” in the mantle?
- # *Exploration, targeting systems* (cf. McCuaig et al., Ore Geol. Rev., 2010)
- # *Defining the LAB* (cf. Jones et al., Lithos, 2010)
- # *Geodynamic modeling (buoyancy from tomography models)*
- # *Lithospheric modeling and evolution (TopoEurope, crustal production*
- # *Dynamic topography (as we learned yesterday!!)*
- # *Craton stabilization... etc...etc...etc*





2) The main goals...

- Predictive and *more explicative* models at regional scales ($\sim 1000 \times 1000$ km) and depths < 600 km.
- How much can we really extract from high-resolution geophysical datasets in terms of thermal and compositional anomalies?
- Realistic uncertainties affecting our predictions

We would like to work within an *internally consistent, multi-observable, probabilistic* inverse framework





Why a *probabilistic* formalism?

Because the problem at hand is probabilistic in nature

Why *multi-observable*?

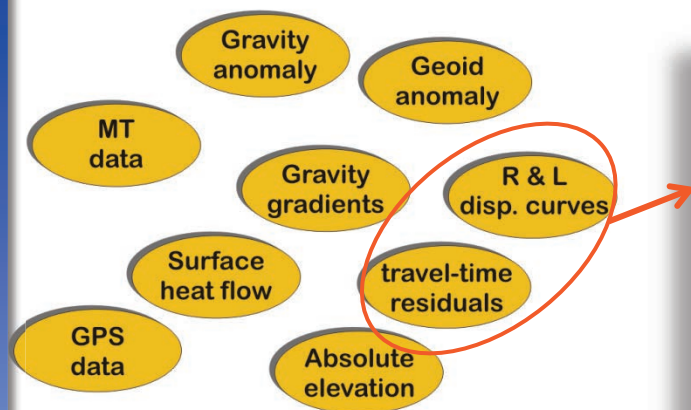
Different observables provide information on different aspects of the problem

Why *internally consistent*?

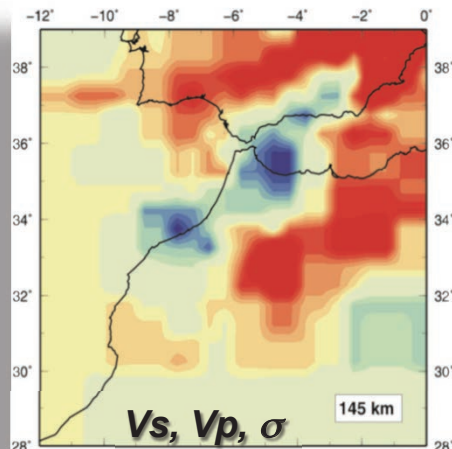
So you cannot tweak parameters as you please to make your model look better!!

3) Multi-Observable Thermochemical Tomography: why and how?

Data

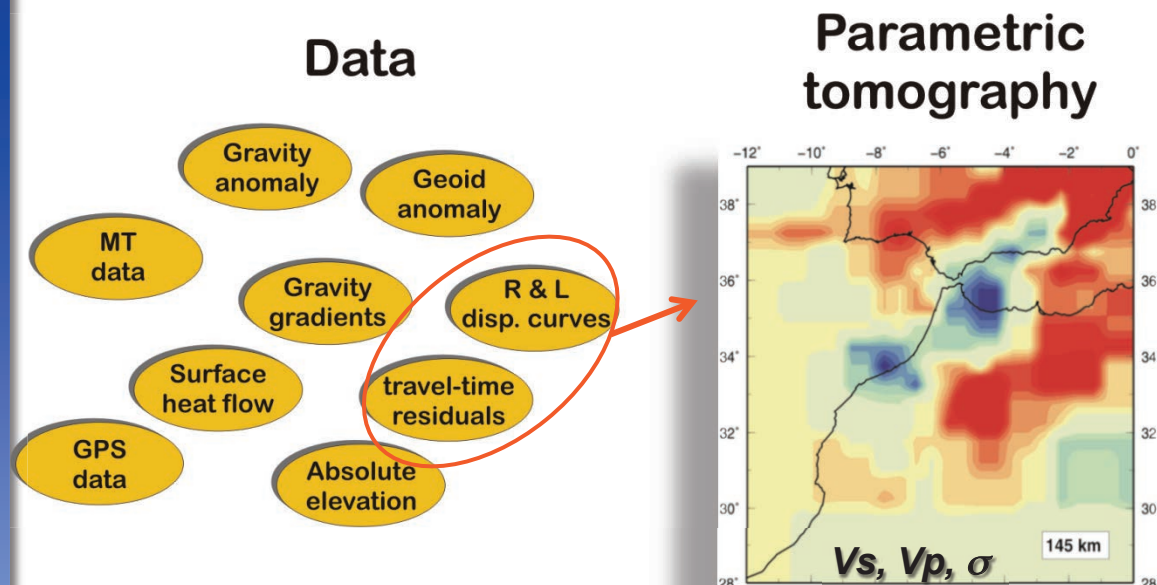


Parametric tomography





3) Multi-Observable Thermochemical Tomography: why and how?



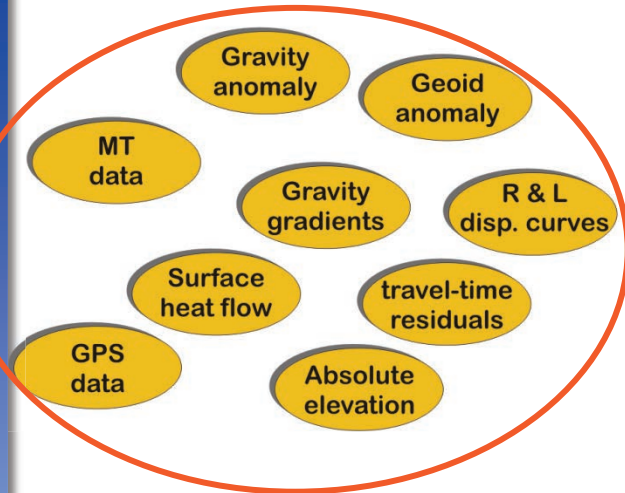
There is abundant complementary information available (e.g. satellite-based gravity, topography, geoid, etc)

Technology capabilities in seismology, mineral physics, geochemical analysis, geodynamics, potential fields, and computing power have reached the required stage of sophistication

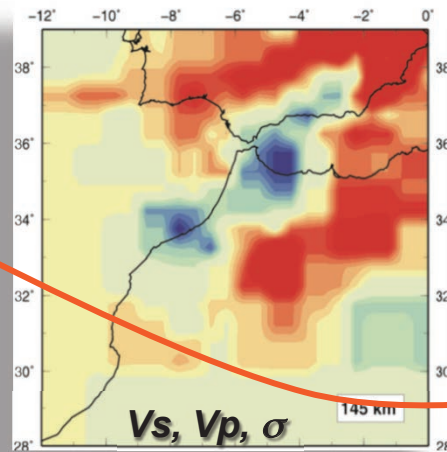
Take the next step: from **parametric tomography** to true **thermochemical multi-observable tomography**

3) Multi-Observable Thermochemical Tomography: why and how?

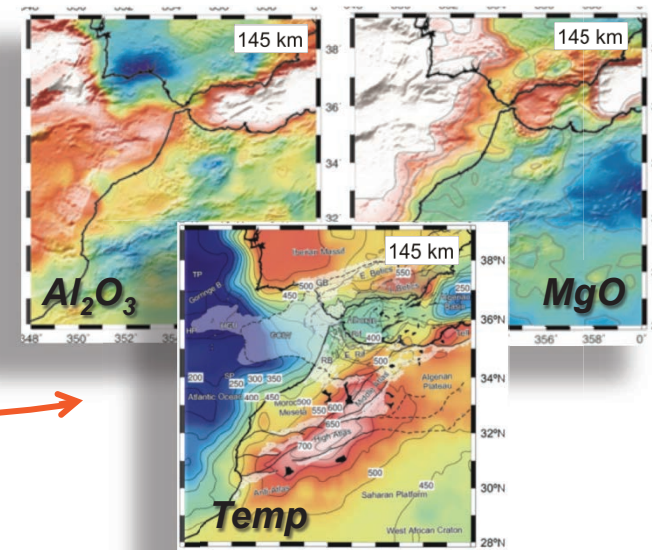
Data



Parametric tomography



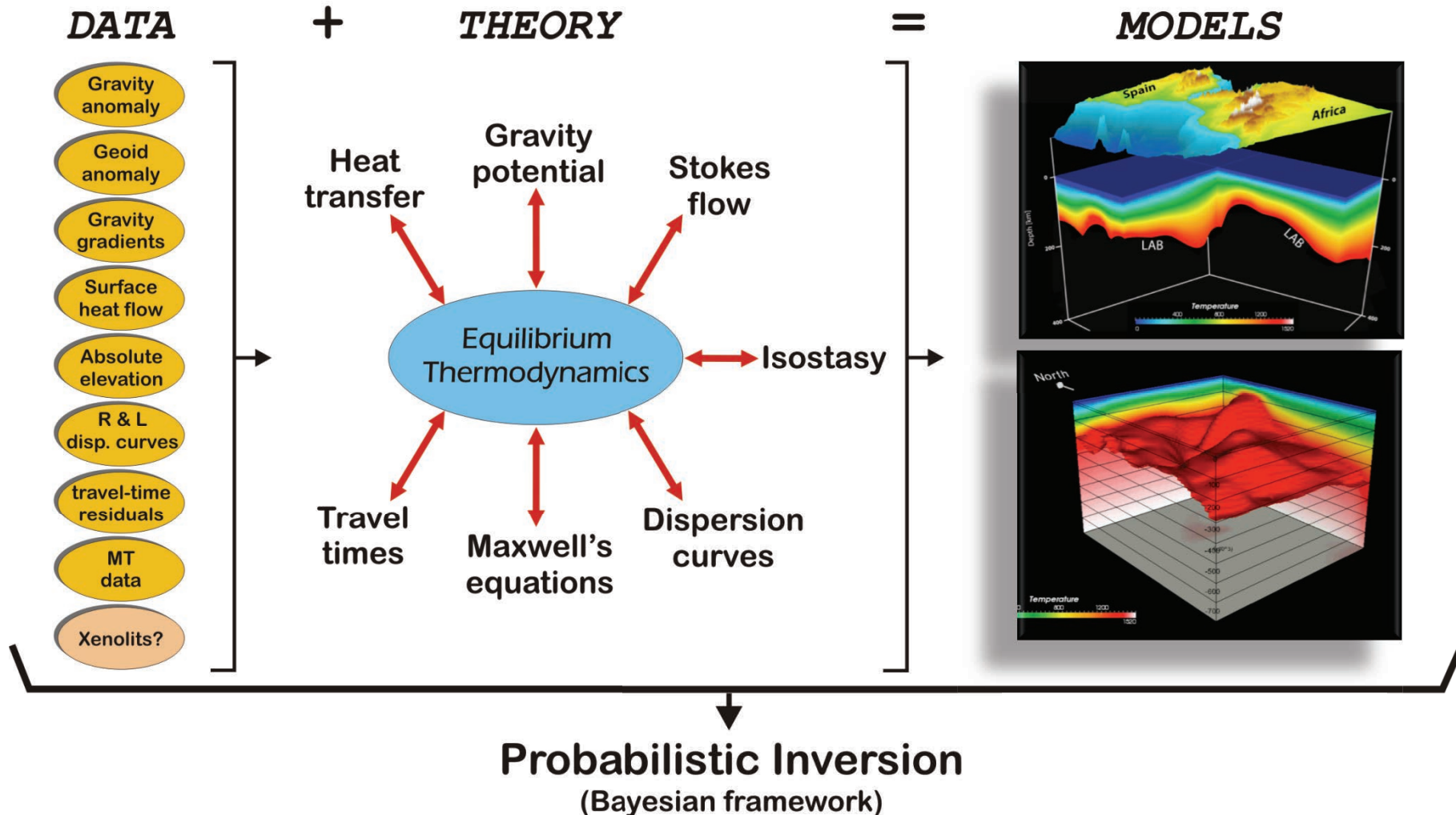
Thermochemical tomography



- # There is abundant complementary information available (e.g. satellite-based gravity, topography, geoid, etc)
- # Technology capabilities in seismology, mineral physics, geochemical analysis, geodynamics, potential fields, and computing power have reached the required stage of sophistication
- # *This is not a methodological competition, but rather a collaborative step forward in methodology*

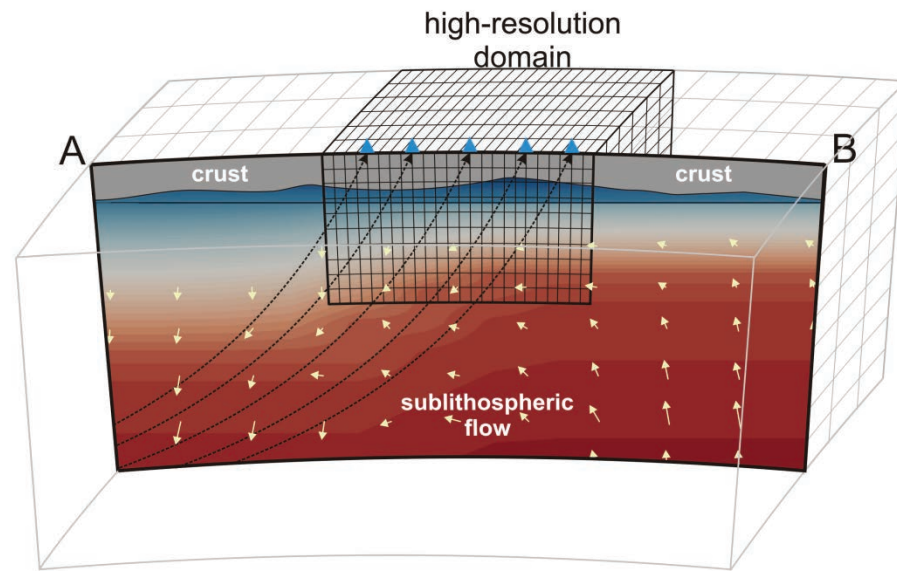
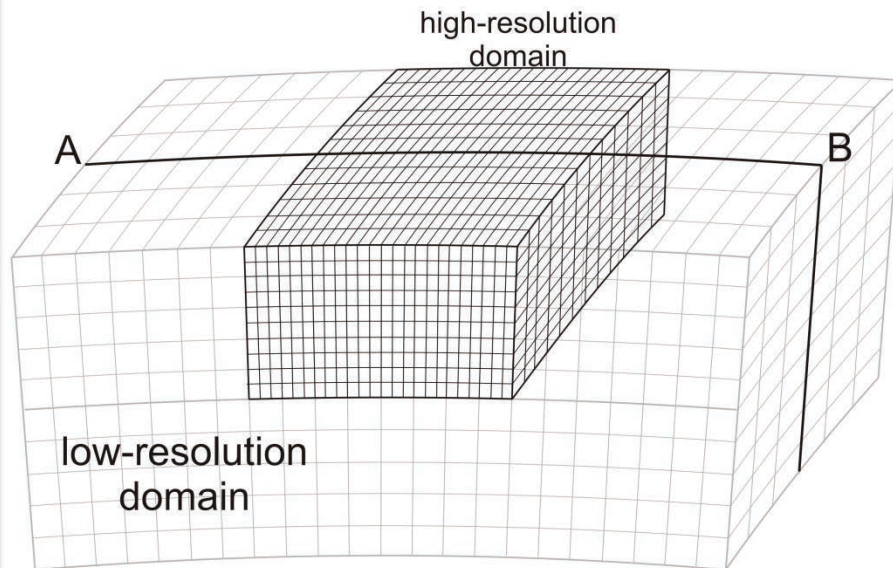
3) Multi-Observable Thermochemical Tomography: why and how?

LitMod_4INV



3) Multi-Observable Thermochemical Tomography: why and how?

Body wave tomography and mantle flow... tough problems



The secret is SIMPLE! ... not to solve a very complicated problem every time we draw a new model, but every n models ... but check convergence!

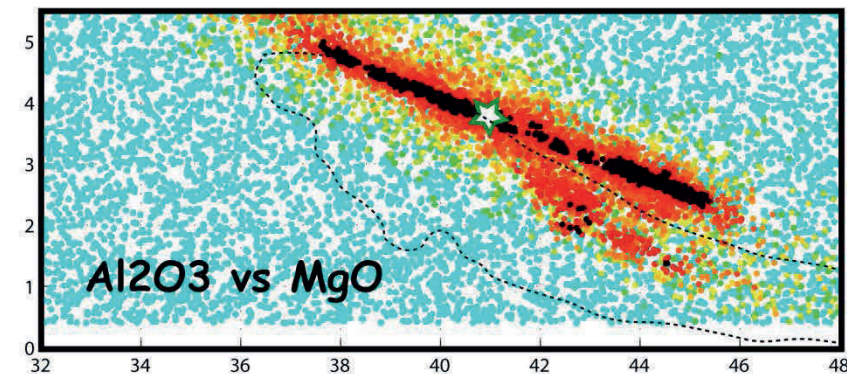
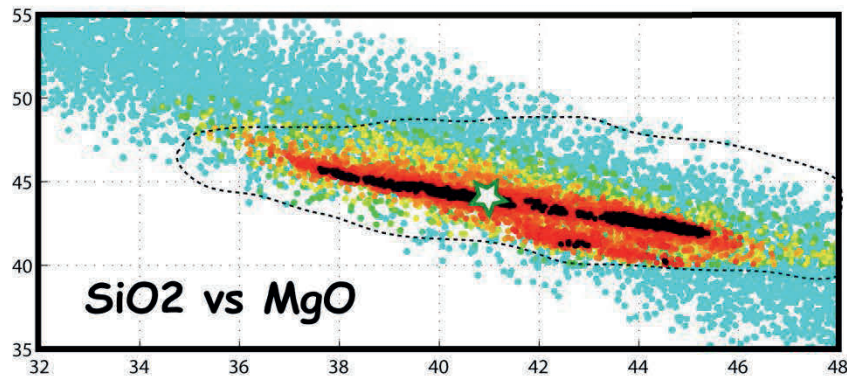
For BWT we're using teleseismic data and a modified version of the Fast Marching Method (Rawlinson and Sambridge, 2005) to compute synthetic travel-time residuals

For the Stokes' flow we are testing a new kind of ultra-fast Stokes solvers based on FEM (w/ G. Rozza & A. Patera, MIT)

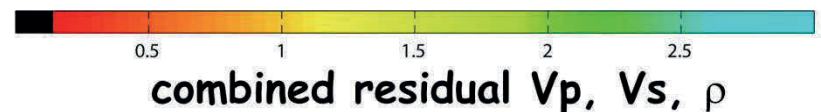
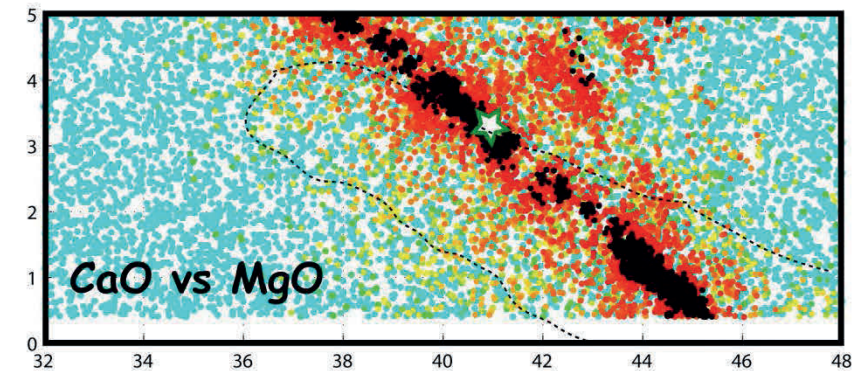
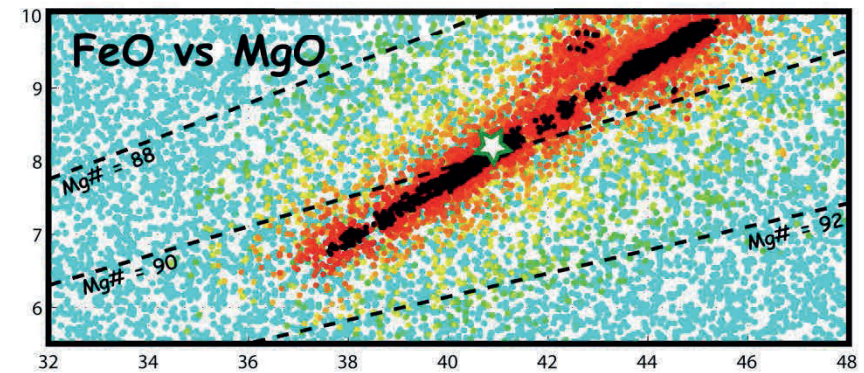
3) Multi-Observable Thermochemical Tomography: why and how?

Non-unique solutions in compositional space

Let's take a "target" composition at
 $T=900^{\circ}\text{C}$ and $P=1.2\text{ GPa}$

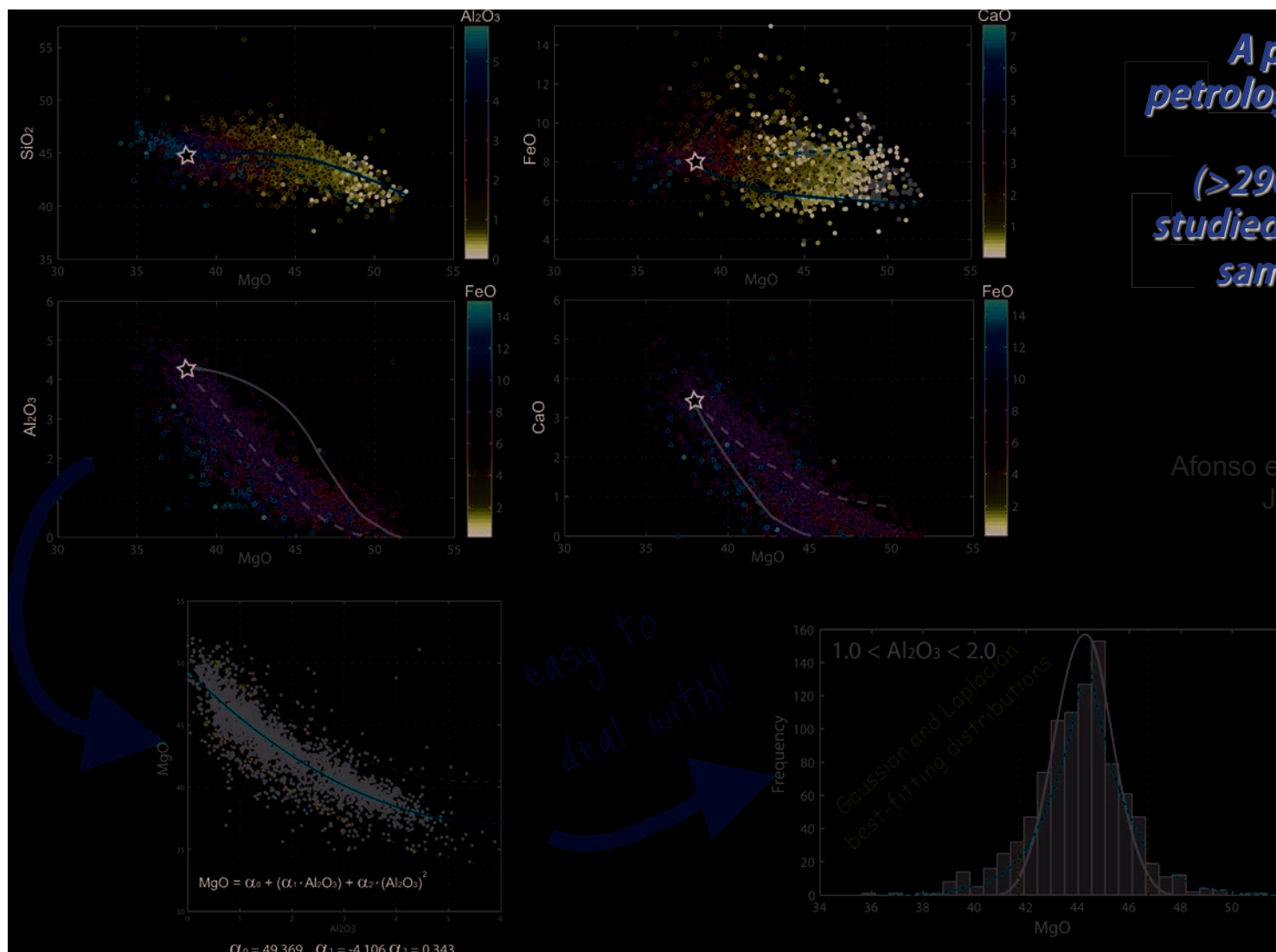


Acceptable models (NA) without a priori information



3) Multi-Observable Thermochemical Tomography: why and how?

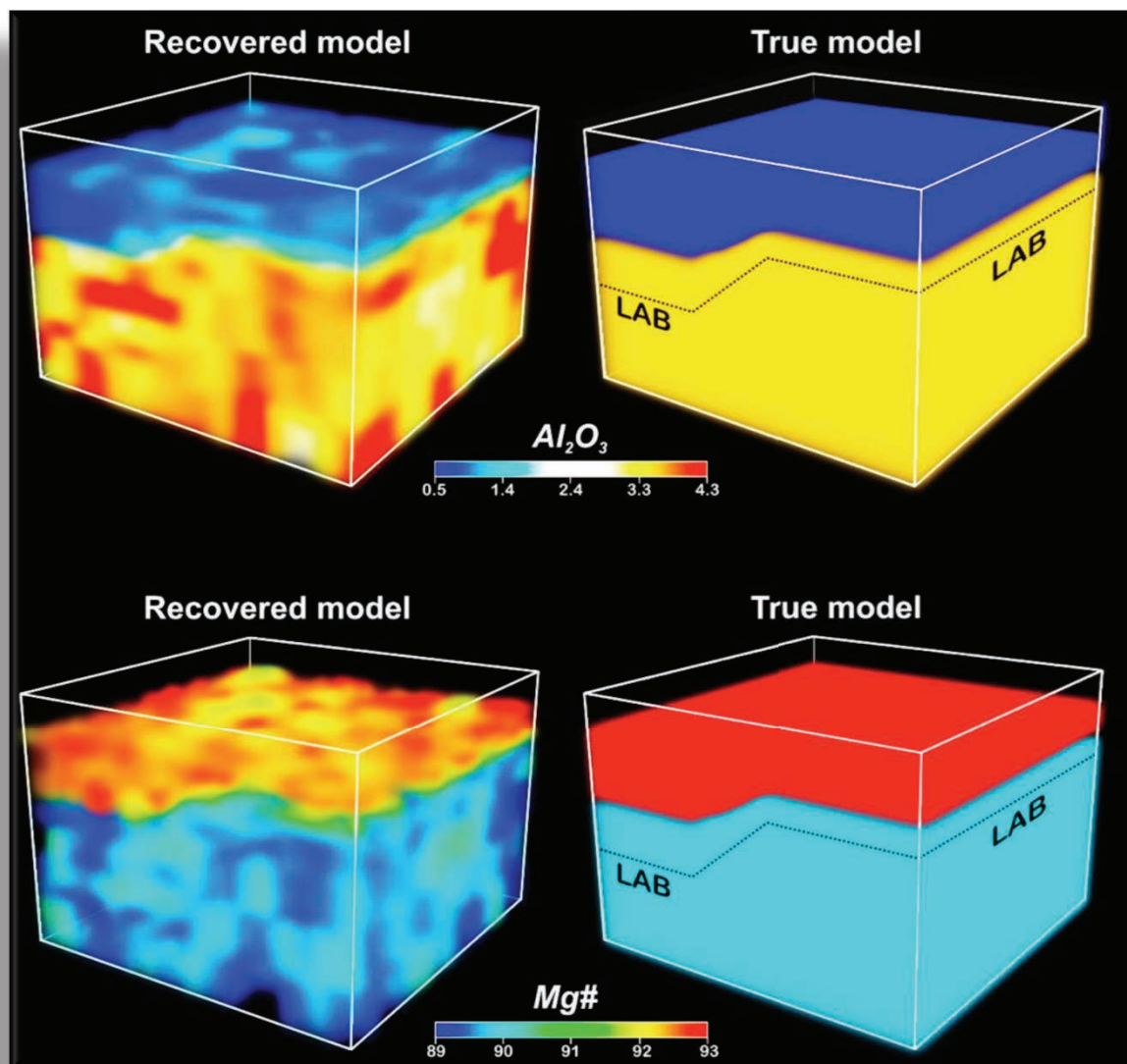
Non-unique solutions in compositional space





3) Multi-Observable Thermochemical Tomography: why and how?

A quick synthetic example:



Shown recovered model is the ML only

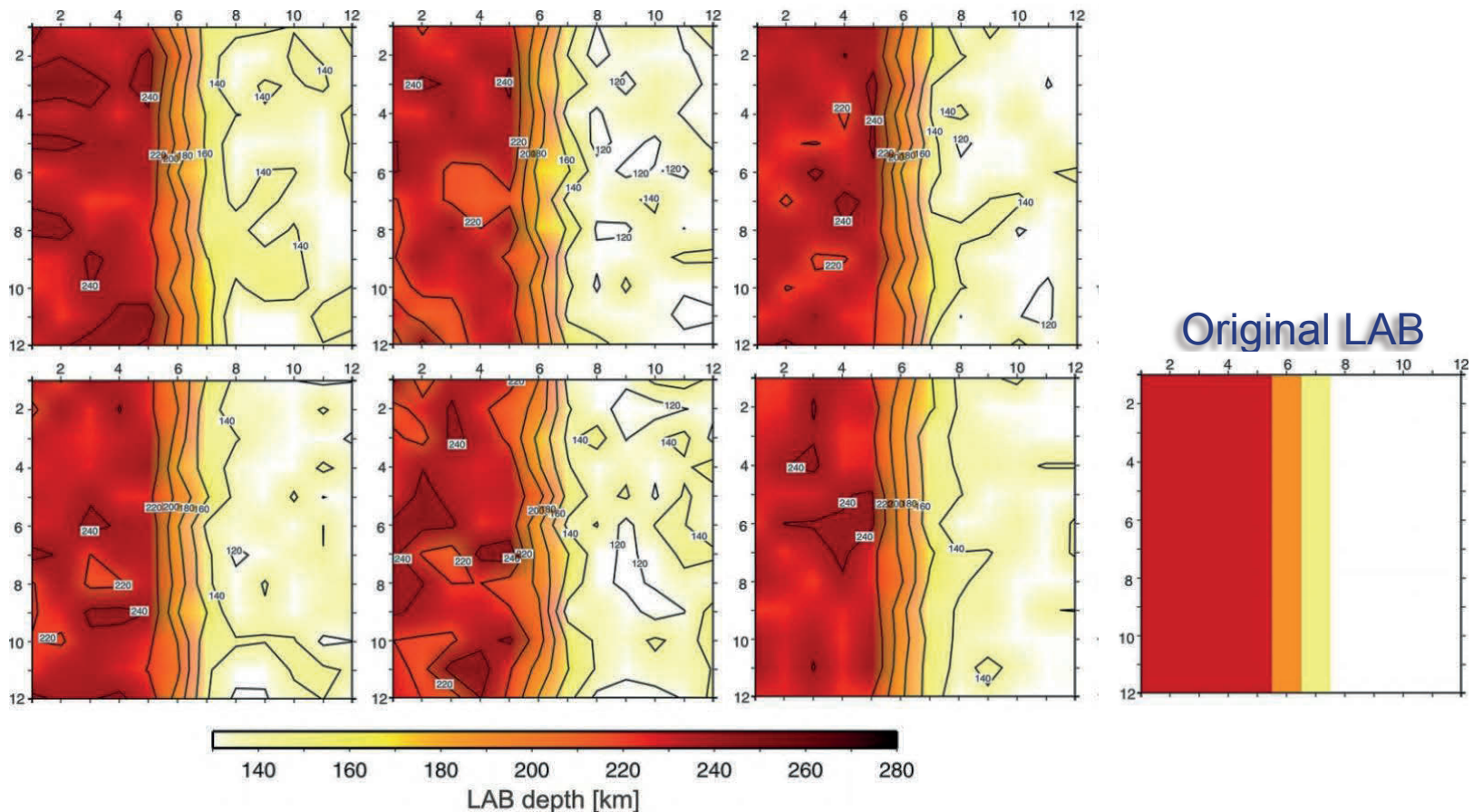
Note the intrinsic variability associated with a unique model

Averaging the PDF would reduce the variability but affects the absolute amplitudes as well

3) Multi-Observable Thermochemical Tomography: why and how?

Results for LAB geometry

Six models taken randomly from within 1 STD of the total posterior PDF





Conclusions

- # *Thermodynamically-constrained multi-observable probabilistic inversions* are particularly well suited for providing reliable estimates of T and C in the upper mantle
- # This approach overcomes or minimizes most of the problems affecting more traditional inversion schemes when applied to the current problem
- # Thermochemical multi-observable tomography is a reality... a computationally expensive one, but we've got supercomputers!
- # Compositional heterogeneities in the sublithospheric mantle... what do they mean?

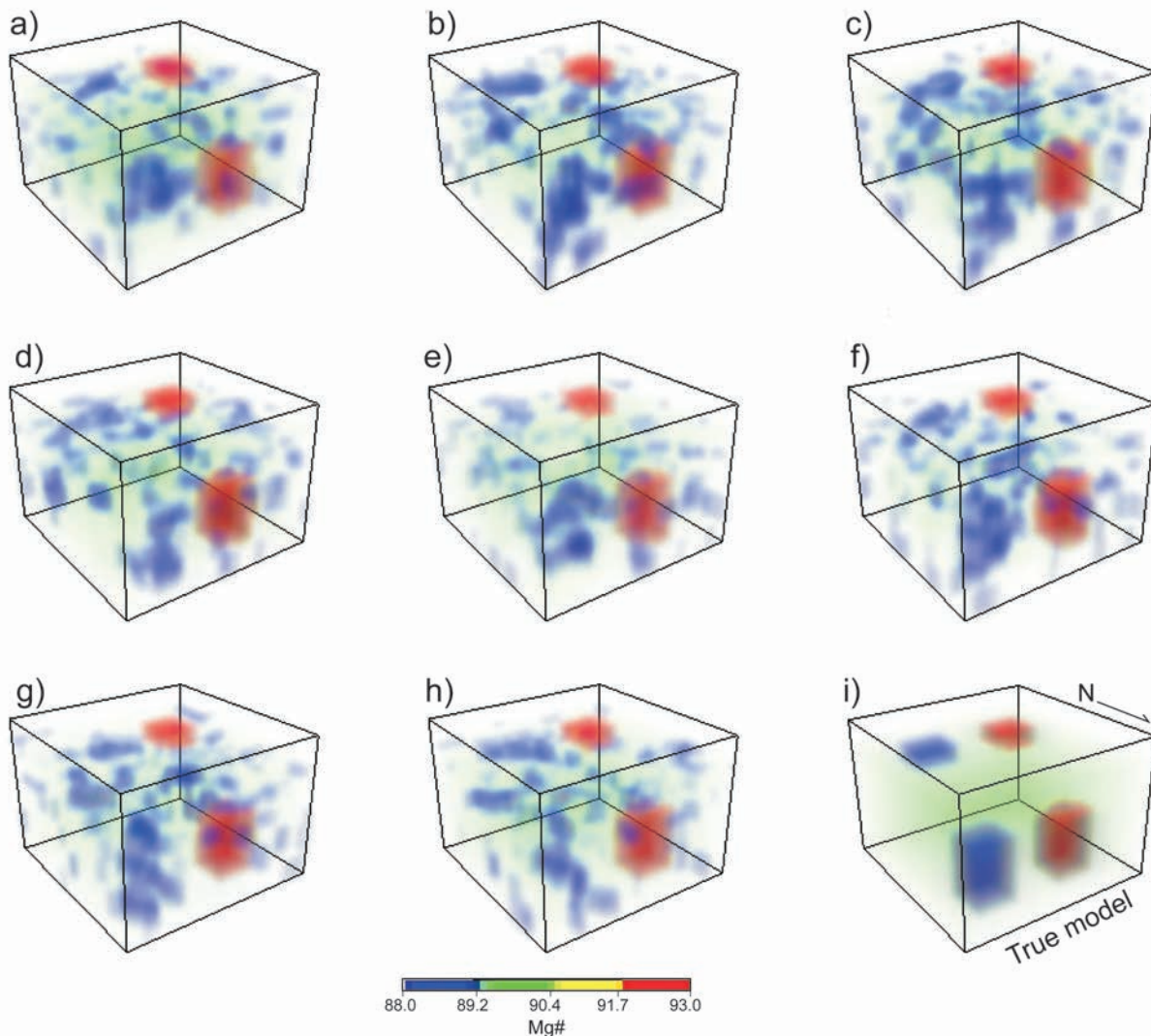


4) A synthetic example

Results for Mg#

Mean models of 8
random ensembles
with 500 samples
each taken from the
total posterior

**Note the
“persistent”
features**

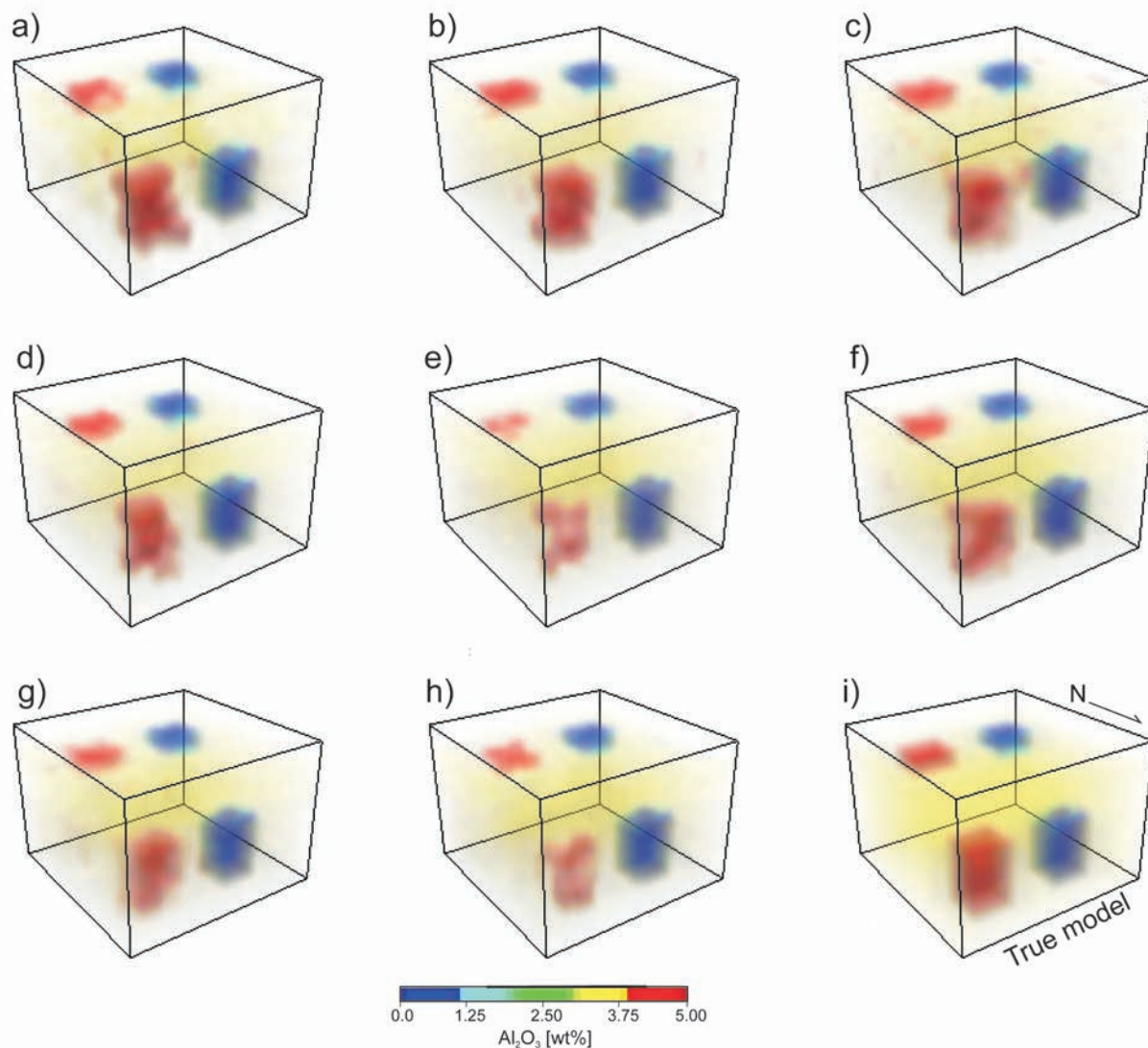


4) A synthetic example

Results for bulk Al_2O_3

Mean models of 8
random ensembles
with 500 samples
each taken from the
total posterior

**Note the
“persistent”
features**



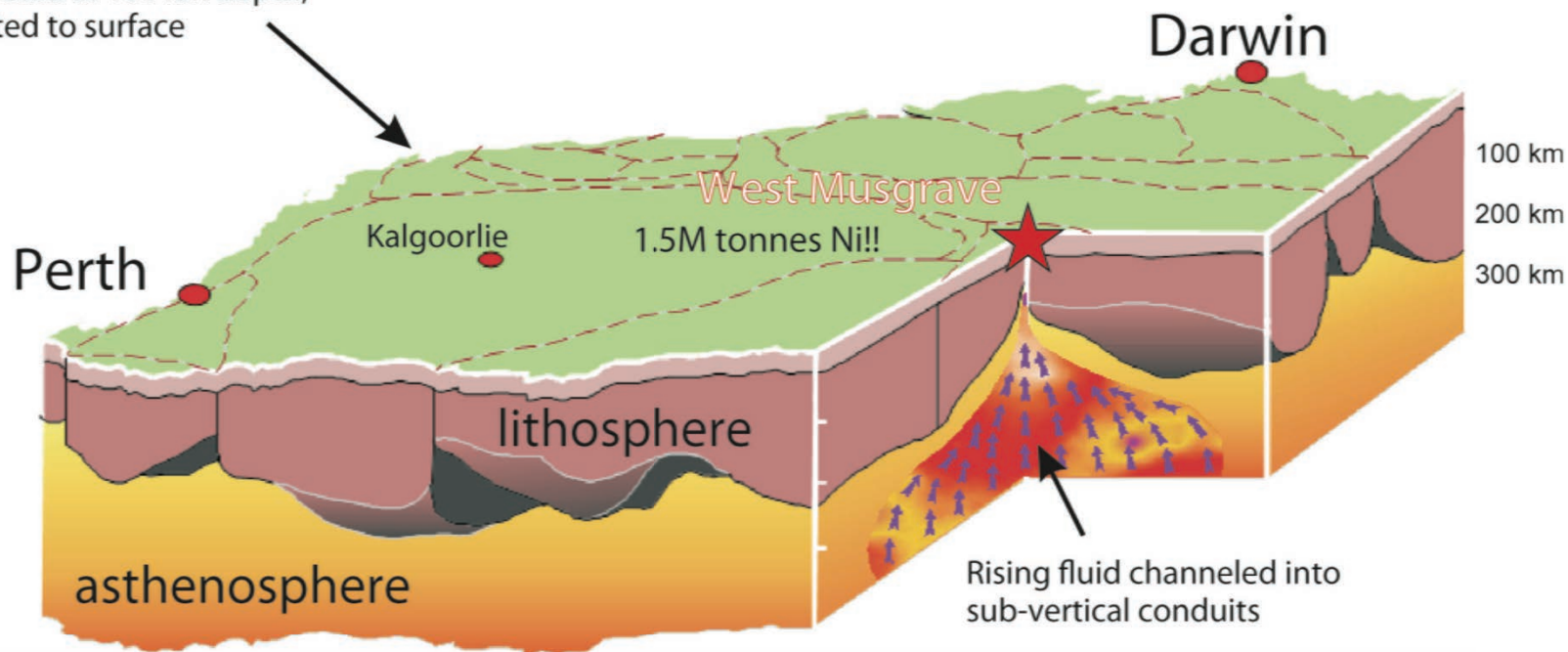


1) An interesting problem...

- What is the nature of the heterogeneity “observed” in the mantle?

Exploration, targeting systems (cf. McCuaig et al., Ore Geol. Rev., 2010)

Upper lithospheric domains
interpreted at 100 km depth,
projected to surface





1) *The main goals...*

Why a *probabilistic* formalism?

Because the problem at hand is probabilistic in nature

Why *multi-observable*?

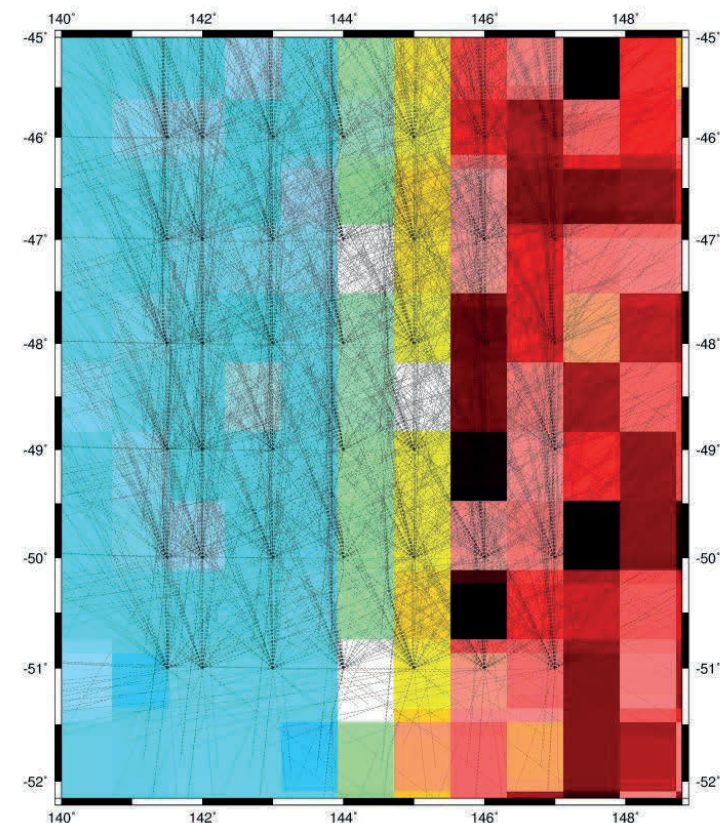
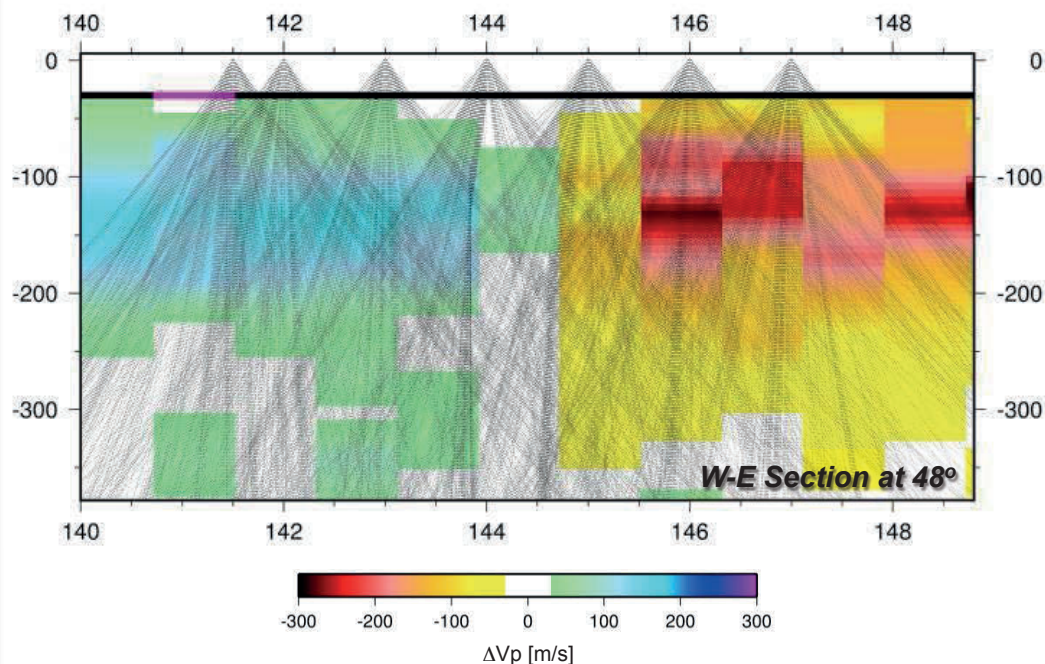
Different observables provide information on different aspects of the problem

Why *internally consistent*?

So you cannot tweak parameters as you please to make your model look better!!

5) Another synthetic example

seismic tomography from posterior

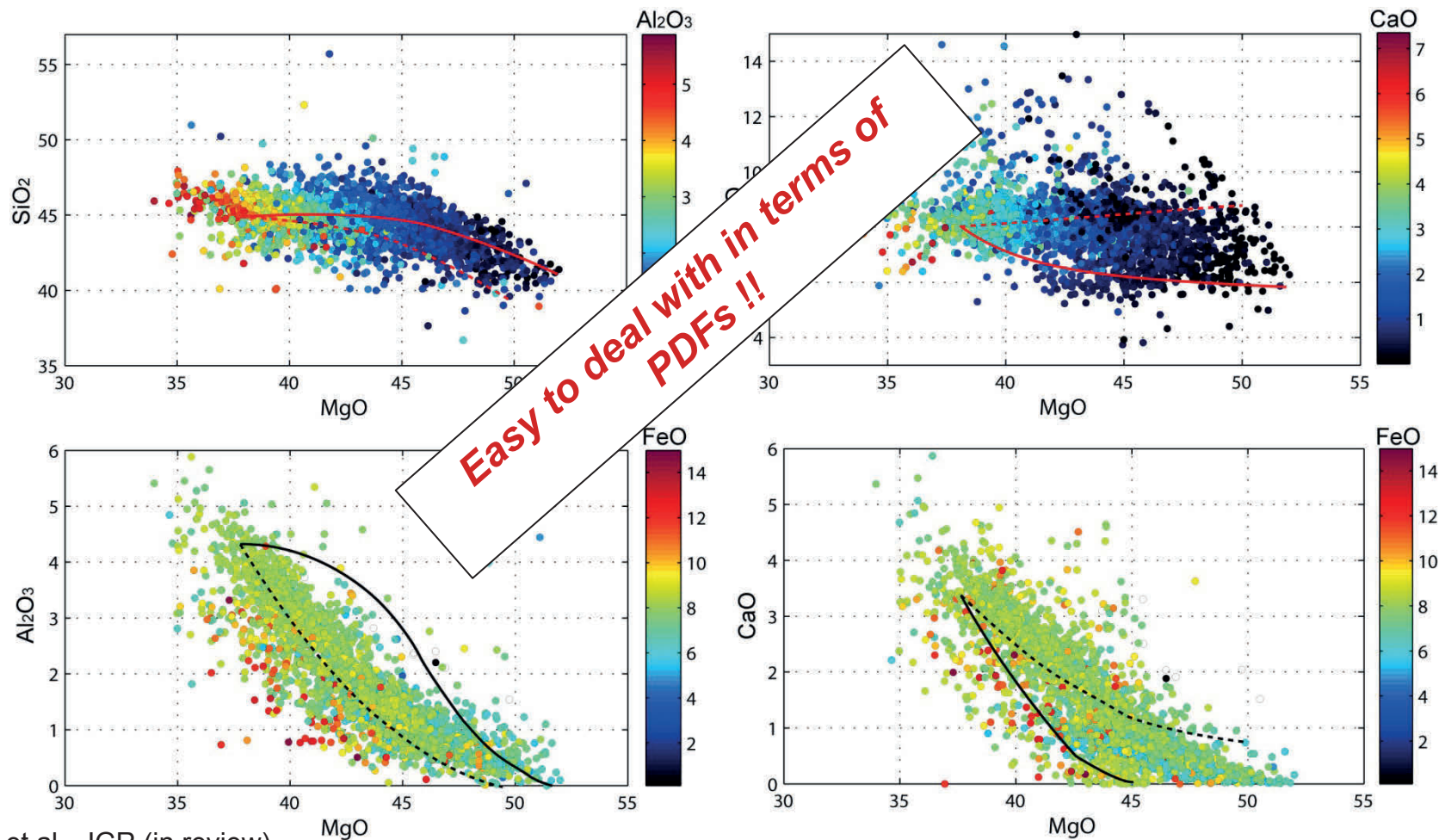


- Our test have been performed with a teleseismic dataset composed of 70 distant sources (from the EVA Array, Victoria) with mostly *P* and *PKiKP* phases (and a few *Pp* and *ScP*)
- The receivers array (42 stations) is synthetic.

3) Our (still preliminary) approach

The *a priori* petrological data

Over 3000 well studied mantle samples





3) How do we do it?

- a) Different observables are sensitive to different chemical-physical properties and depth ranges
- b) Each method is designed to specific chemical-physical properties or perturbations of these properties
- c) All thermophysical properties of interest ultimately depend on **T**, **P**, **C**

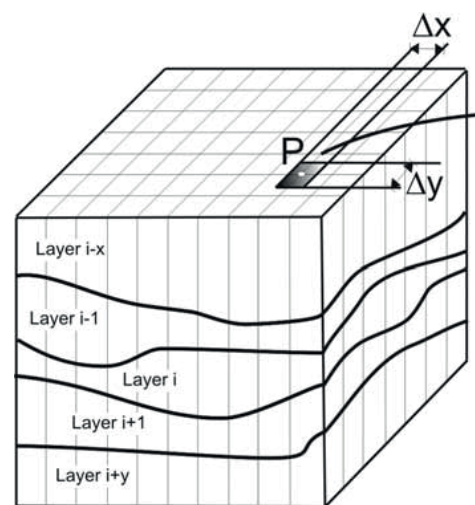
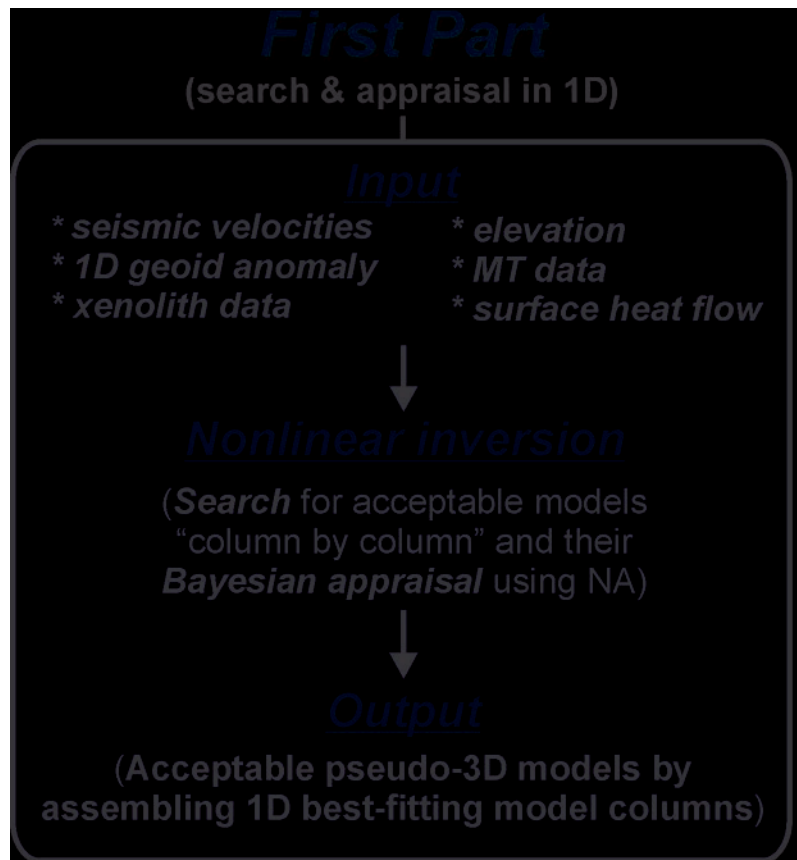
$$dG = V d\mathbf{P} - S d\mathbf{T} + \sum_i \mu_i d\mathbf{n}_i$$

$$V = \left(\frac{\partial G}{\partial P} \right)_T \quad V\alpha = - \left(\frac{\partial S}{\partial P} \right)_T = \left(\frac{\partial^2 G}{\partial P \partial T} \right) \quad C_P = -T \left(\frac{\partial^2 G}{\partial T^2} \right)_P \quad c_{ijkl} = \frac{1}{V} \left(\frac{\partial^2 G}{\partial S_{ij} \partial S_{kl}} \right)_{P,T}$$

All related to the free energy of the system (in equilibrium)

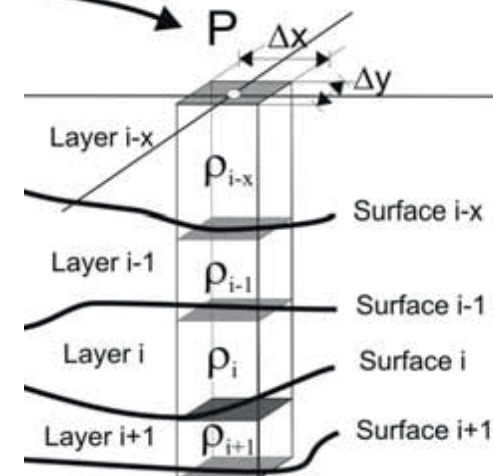
4) Our method

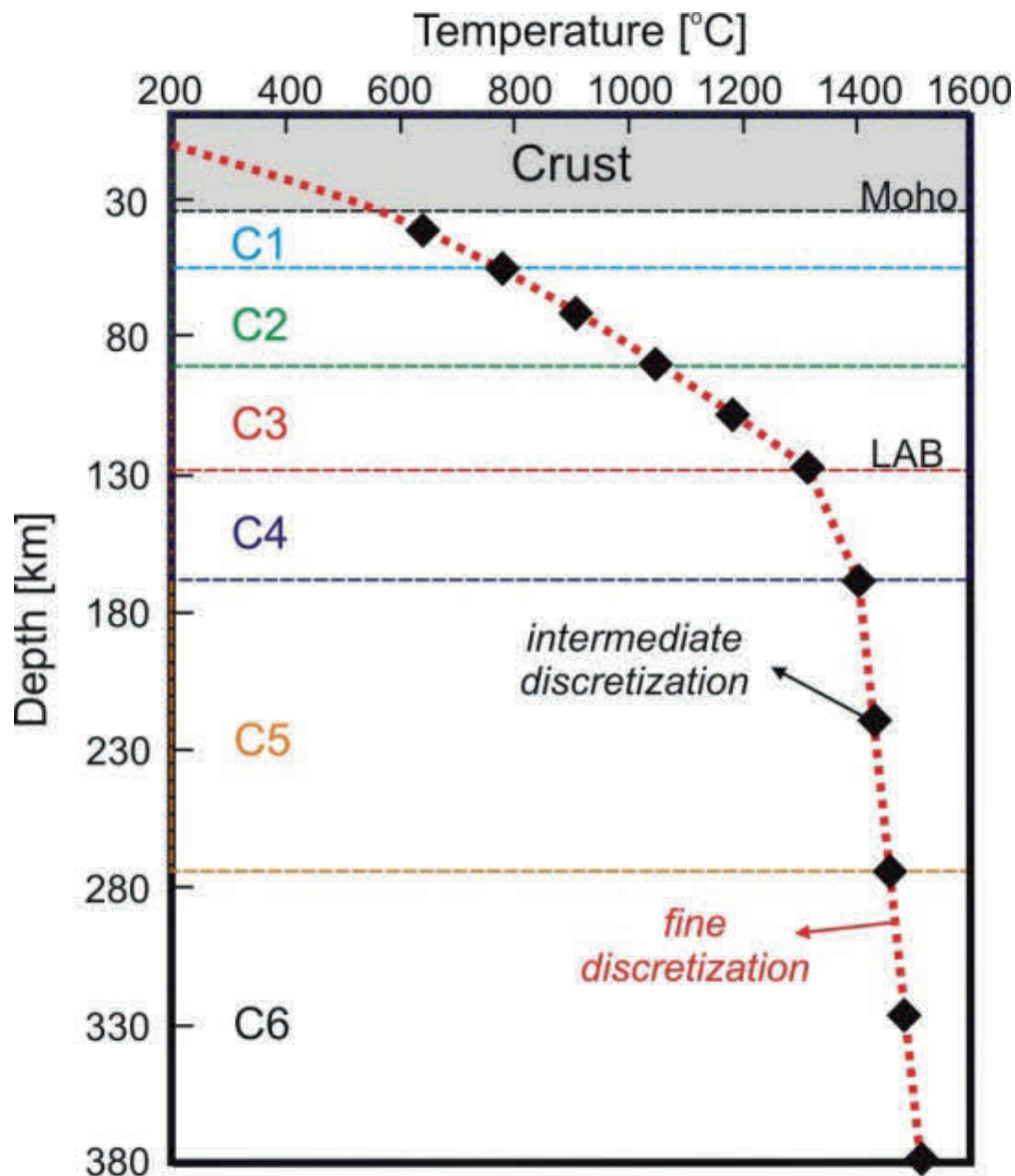
* There are two main parts



3D discretization

Individual columns







3) The main problems...

- i) Nonlinearity of the problem at hand
- ii) Thermodynamic modelling
- iii) Trade-off between T and C in wave speeds
- iv) T effect is much stronger than C effect (i.e. hard to isolate)



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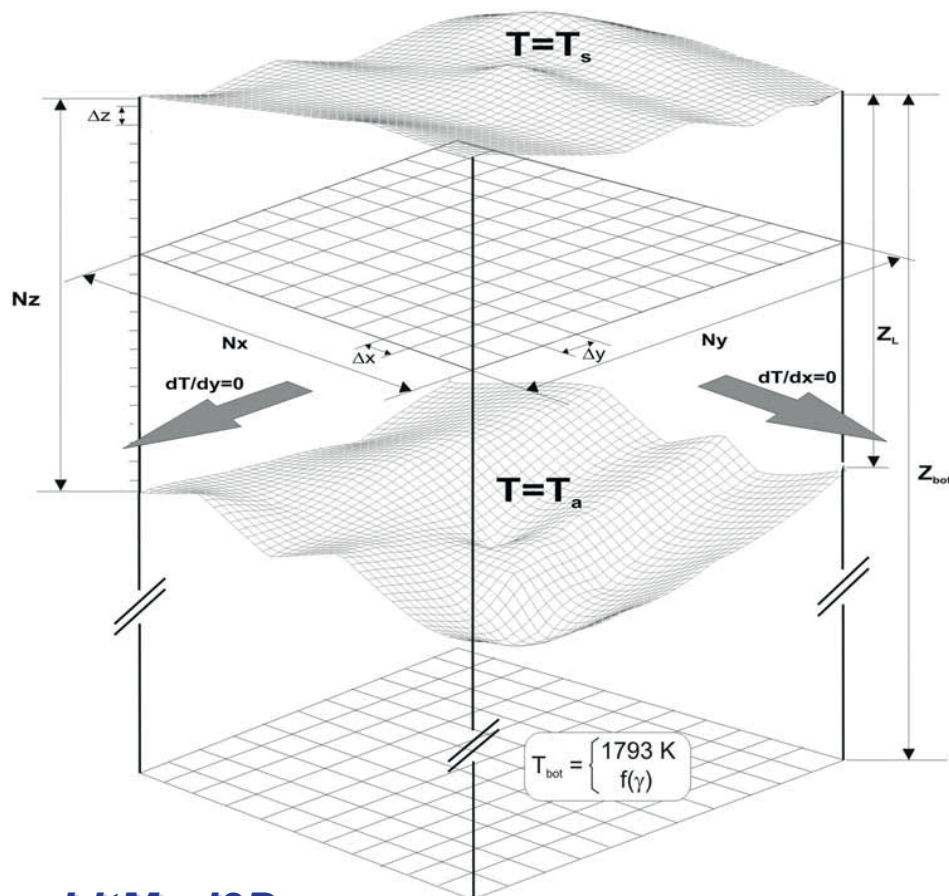
On the V_p/V_s –Mg# correlation in mantle peridotites: Implications for the identification of thermal and compositional anomalies in the upper mantle

Juan Carlos Afonso ^{a,*}, Giorgio Ranalli ^b, Manel Fernàndez ^c, William L. Griffin ^a,
Suzanne Y. O'Reilly ^a, Ulrich Faul ^d



4) Our method

* There are two main parts



LitMod3D

<http://www.eps.mq.edu.au/~jafonso/Software1.htm>

Second Part

(refinement in 3D)

Input

- * seismic velocities
- * 3D geoid anomaly
- * 3D gravity anomalies
- * xenolith data
- * elevation
- * 3D MT data
- * surface heat flow
- * output from part 1

Inversion/Forward

pure forward or non-linear LS inversion or Bayesian adaptive MCMC

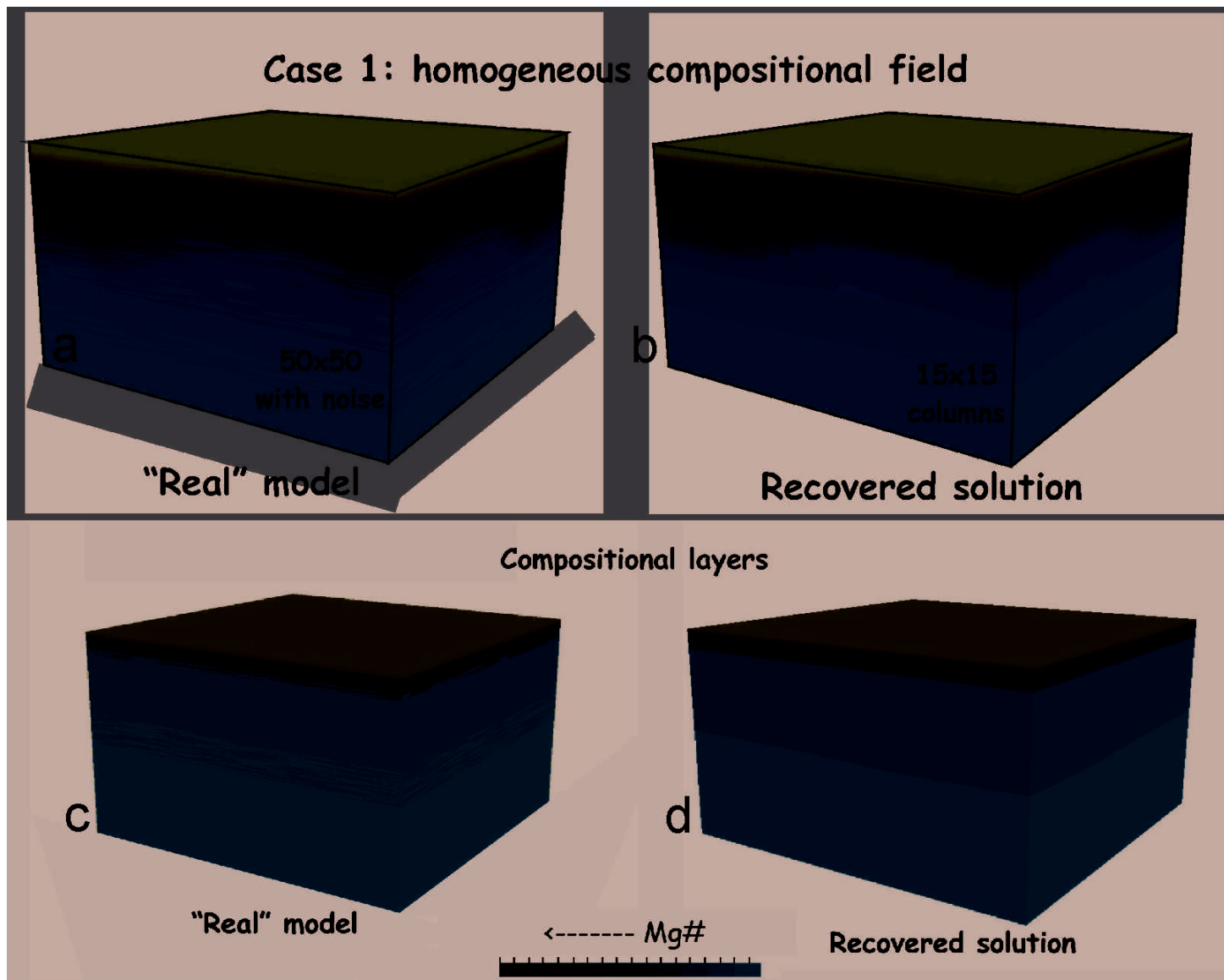
Output

(Real 3D model of the compositional and thermal structure of the lithosphere and upper mantle + uncertainties)



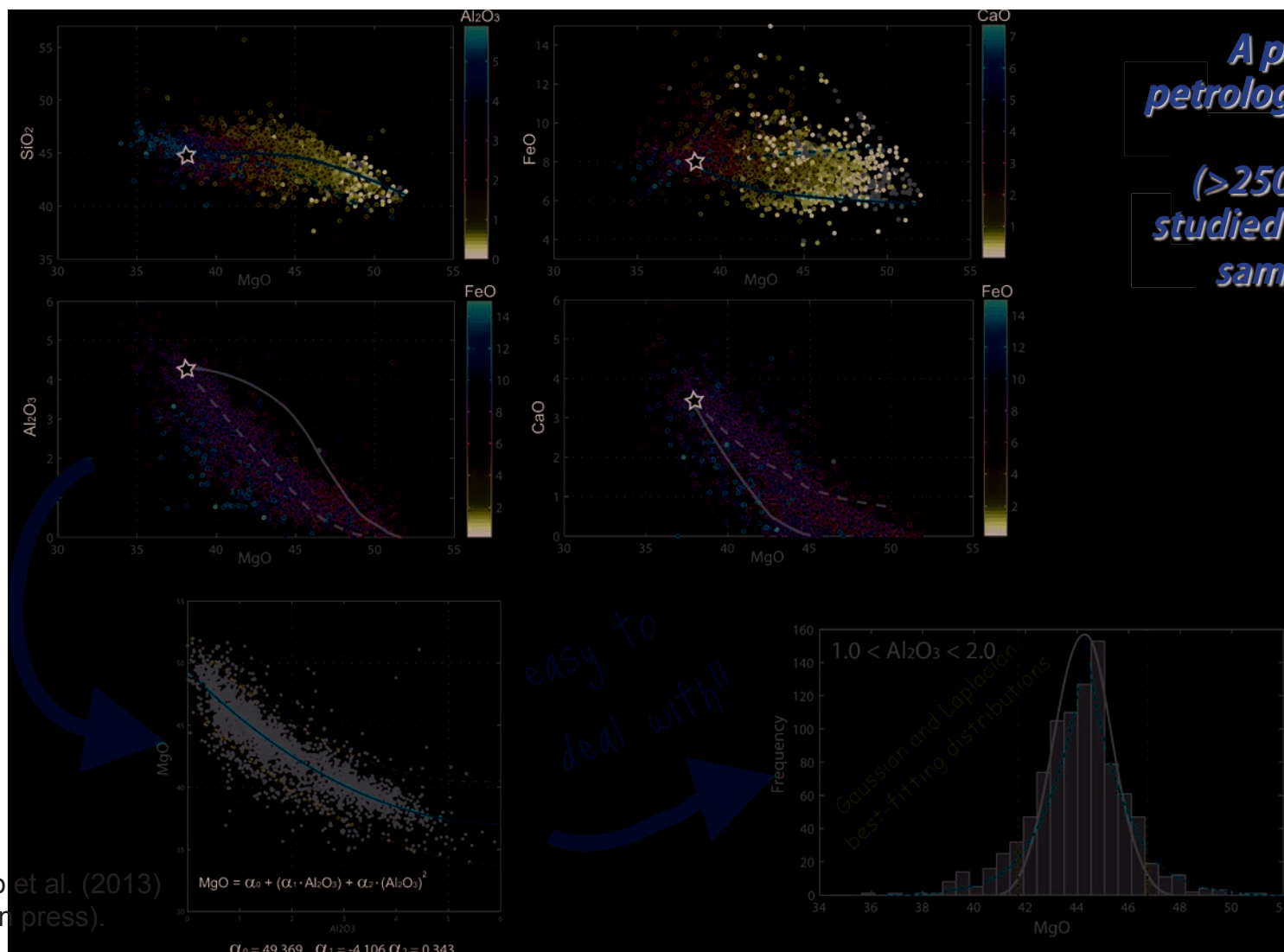
5) Preliminary results (synthetic cases)

*Simulation
run in a 90-
CPU cluster
for ~ 10 days*



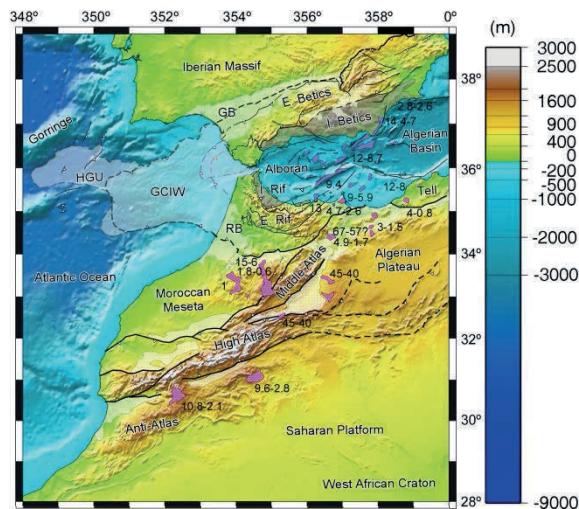
3) The main problems...

Non-unique solution in compositional space



Afonso et al. (2013)
JGR (in press).

5) Preliminary results (real case)

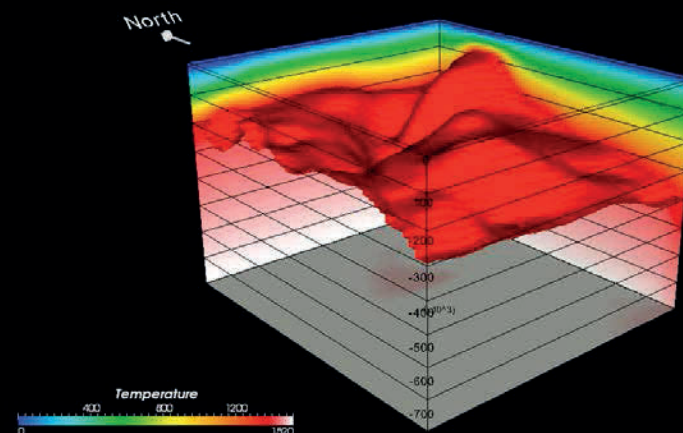
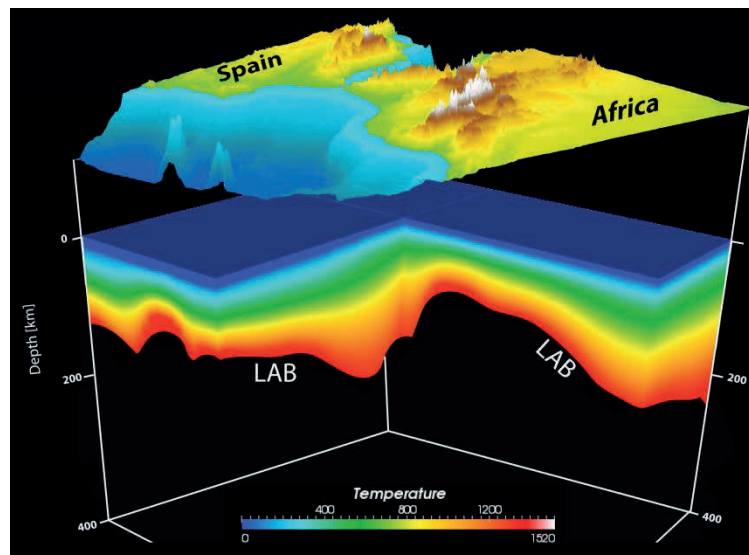


From Fullea et al., 2010, *Lithos*

An example...

the Atlantic-Mediterranean
Transition Region

It's a highly complicated and
interesting area





3) How do we do it?

First Part

(search & appraisal in 1D)

Input

- * Dispersion curves
- * 1D geoid anomaly
- * xenolith data
- * Vp structure
- * elevation
- * MT data
- * surface heat flow



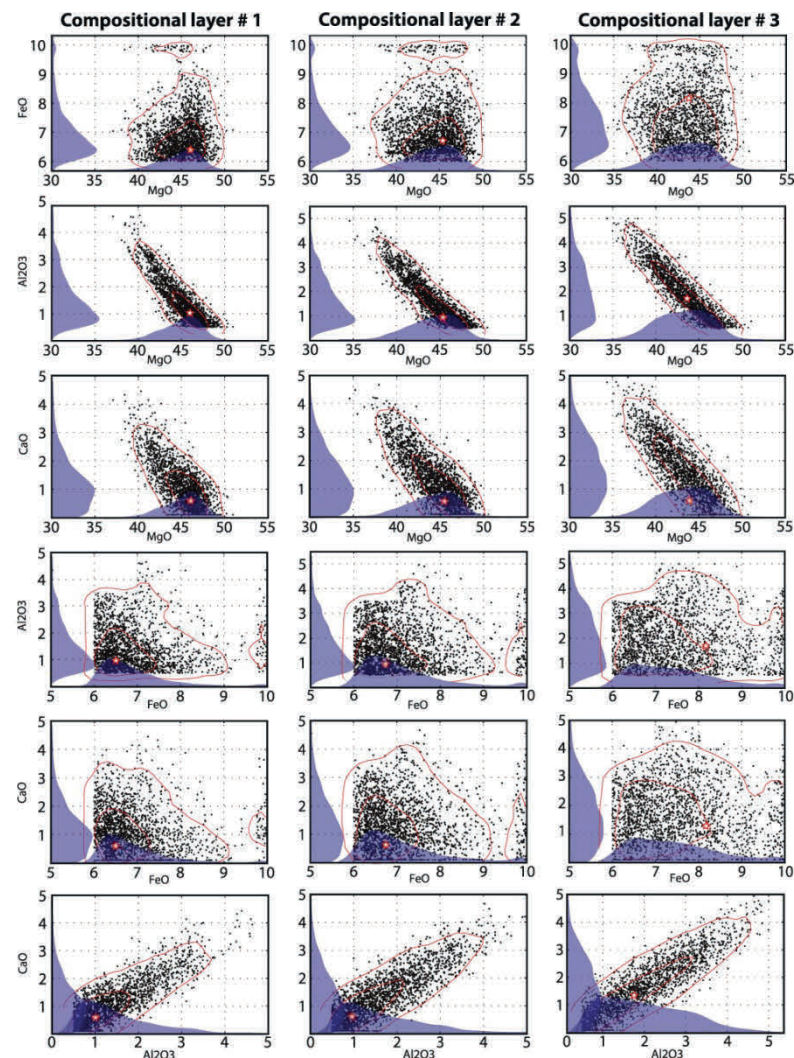
Nonlinear inversion

(**Search** for acceptable models
“column by column” and their
PDFs using NA)

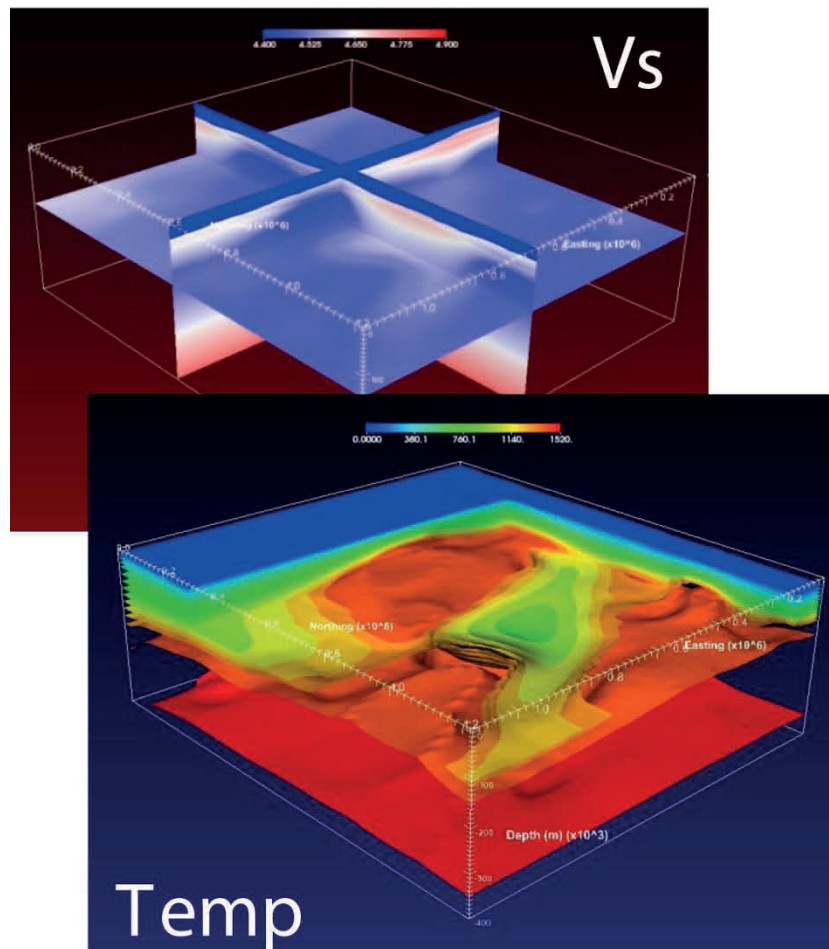


Output

(PDFs for all parameters
in all individual columns)



3) How do we do it?



Afonso et al., in prep.

Second Part (refinement in 3D)

Input

- * Dispersion curves
- * 3D geoid anomaly
- * 3D gravity anomalies
- * xenolith data
- * Vp structure
- * elevation
- * MT data
- * surface heat flow
- * **output from part 1**

Inversion/Forward

pure
forward

or

non-linear
LS inversion

or

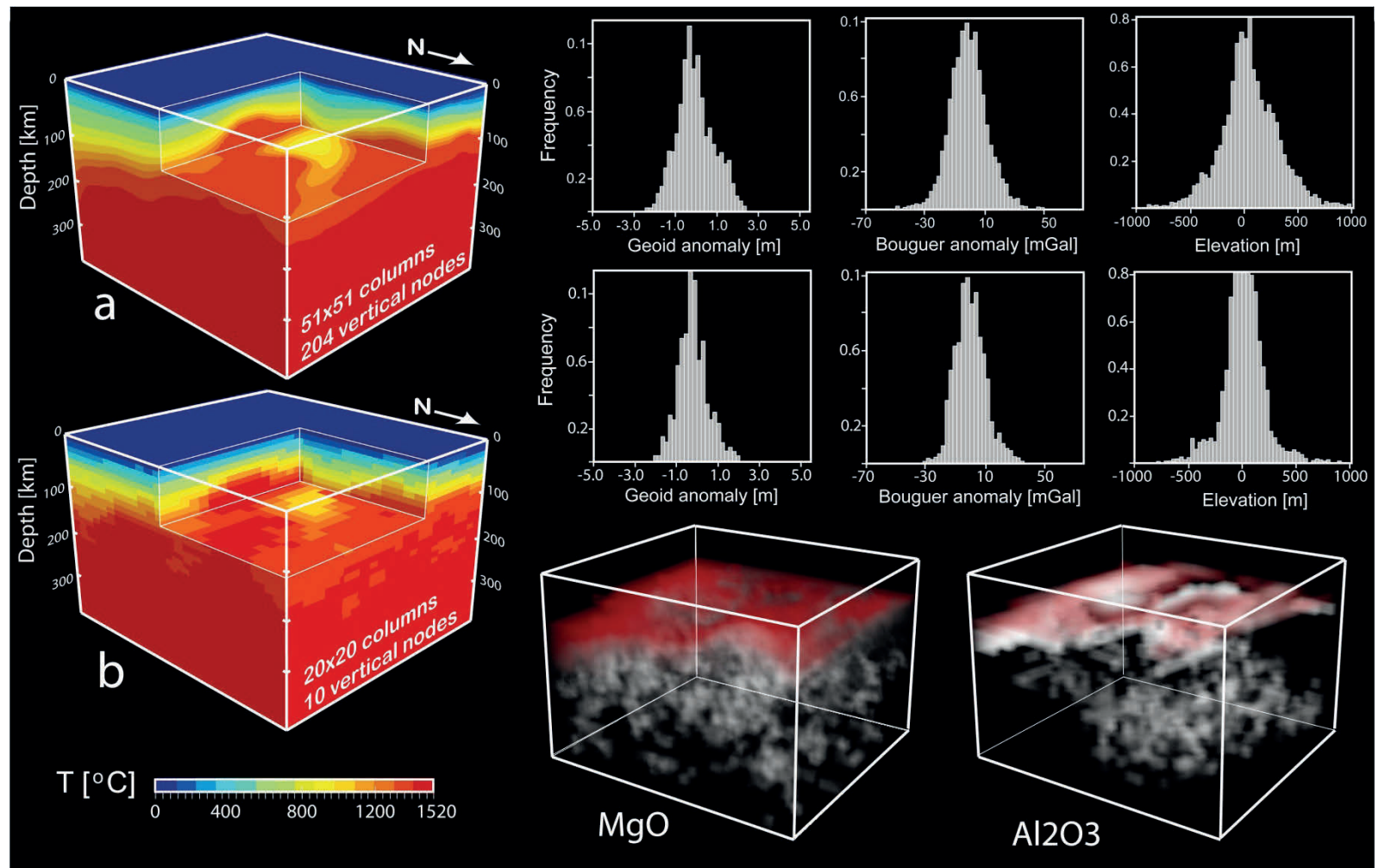
Bayesian
MCMC

Output

(Posterior PDF of the 3D compositional and thermal structure of the lithosphere and sublithospheric upper mantle + uncertainties)

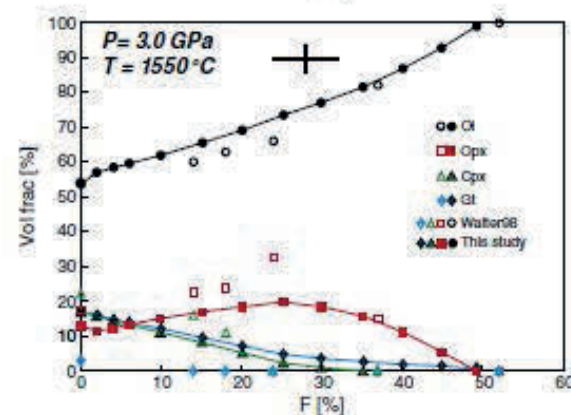
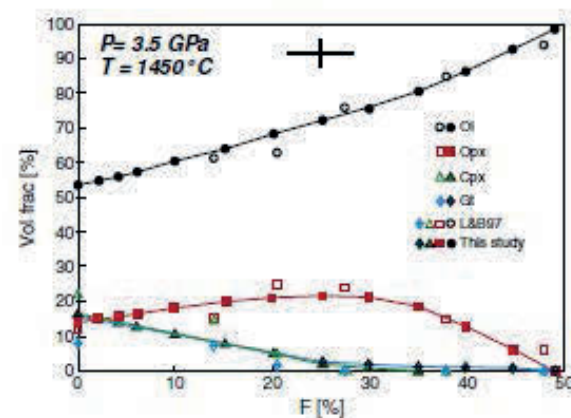
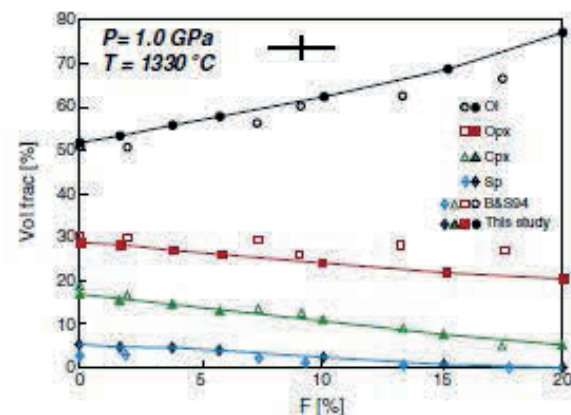
5) Preliminary results

A comparison between the results of Fullea et al. (2010) and our new method



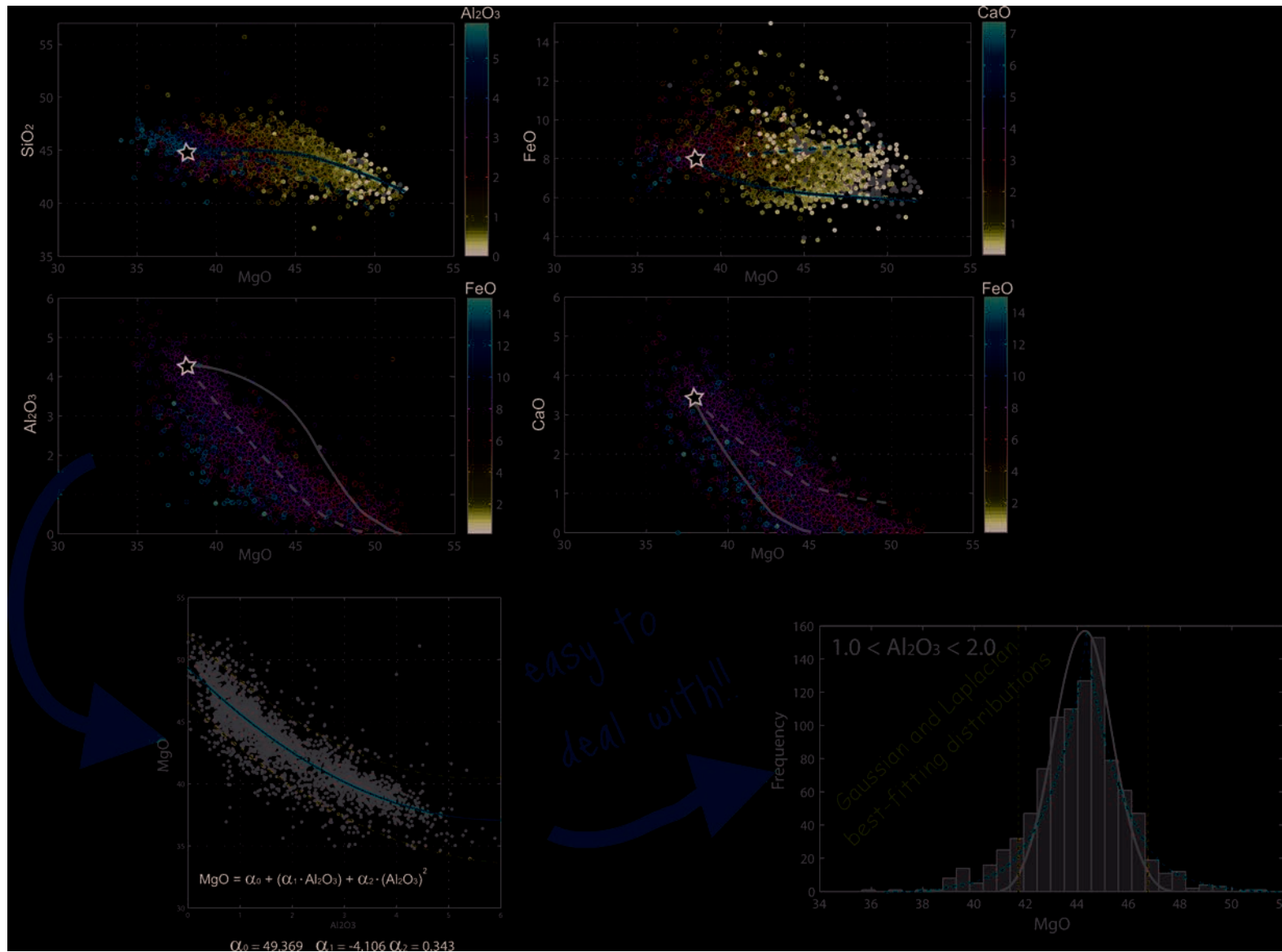


Thermodynamics



From Afonso and Schutt., *Lithos* (2012)

2) How do we do it?





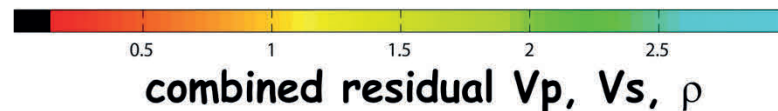
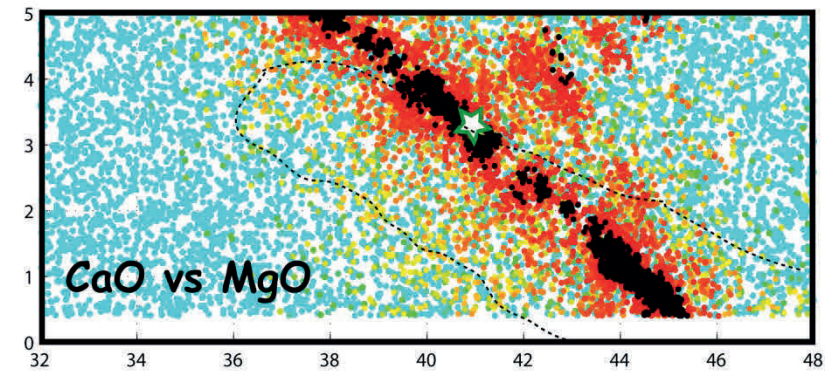
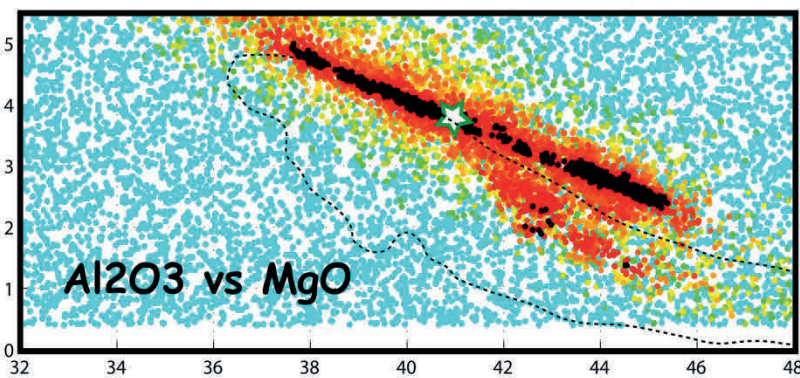
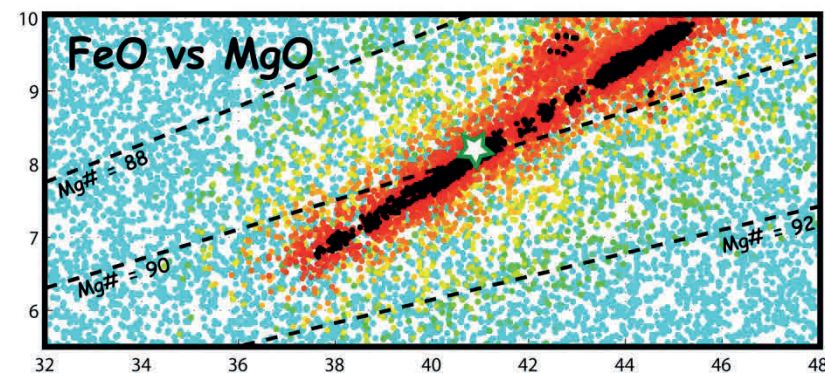
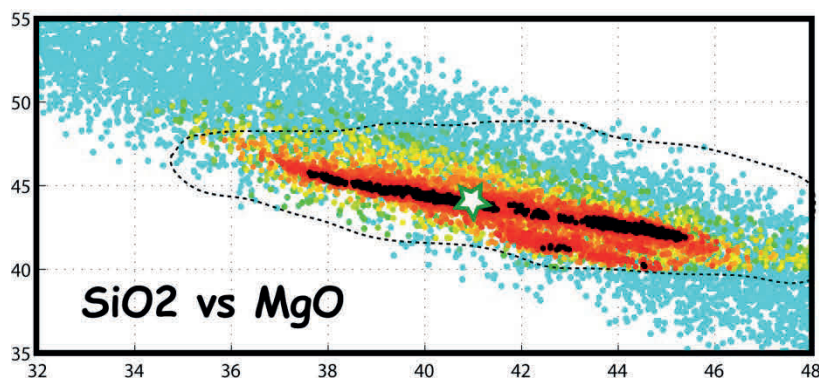
2) *Why bother with this?*

- # Defining the LAB (cf. Jones et al., Lithos, 2010)
- # Exploration, targeting systems (cf. McCuaig et al., Ore Geol. Rev., 2010)
- # Geodynamic modeling (buoyancy from tomography models)
- # Lithospheric modeling and evolution (TopoEurope, crustal production trough time)
- # Craton stabilization... etc...etc...etc

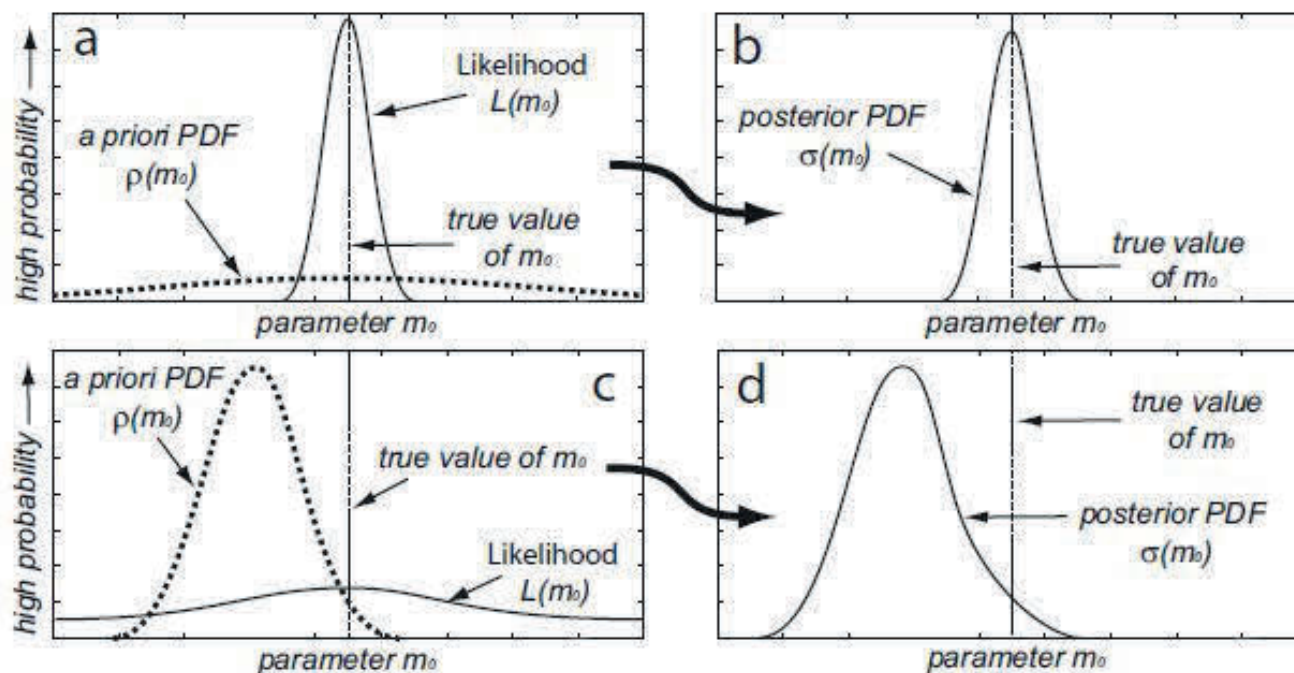
3) Should it work?

Let's take a "target" composition at $T=900\text{ }^{\circ}\text{C}$ and $P=1.2\text{ GPa}$

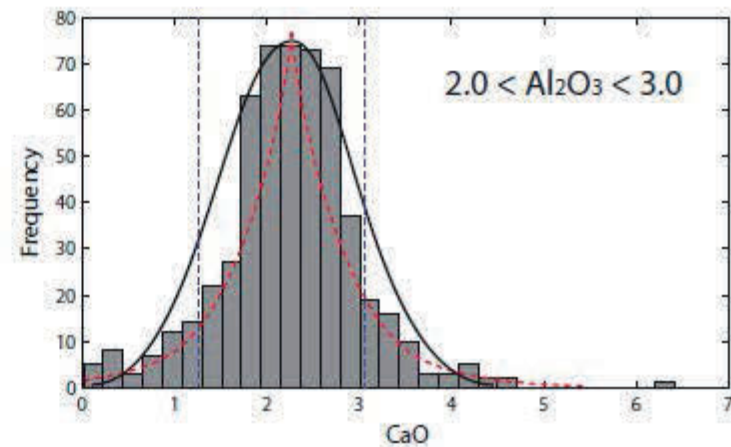
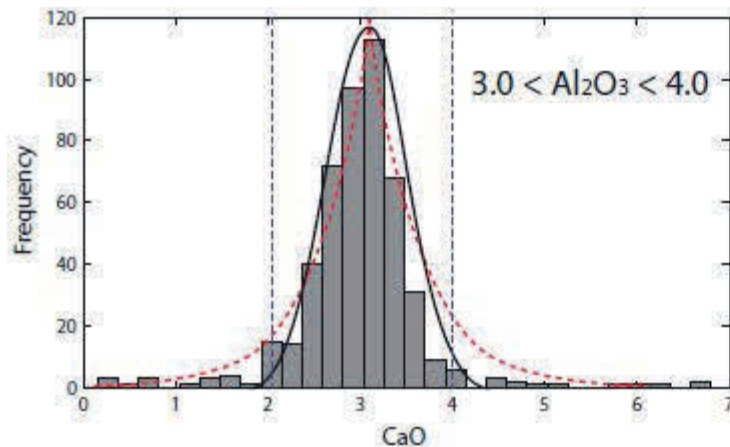
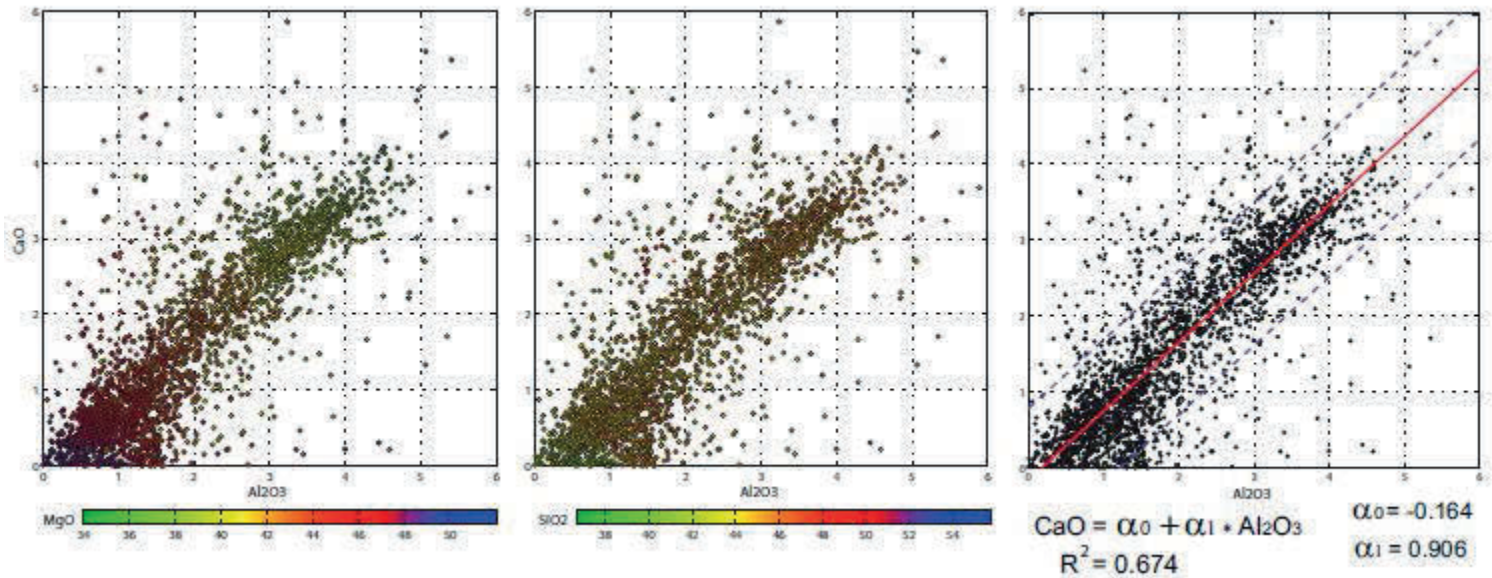
Acceptable models (NA) without any a priori information



$$\sigma(m) = k\rho(m)L(m)$$

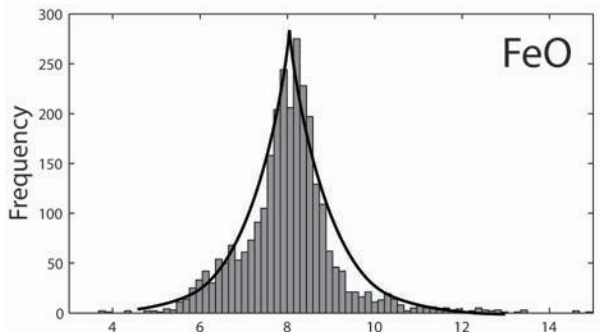
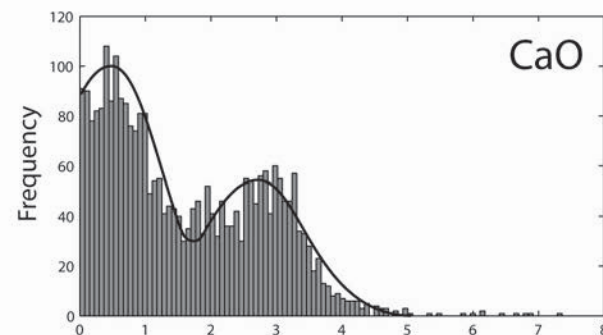
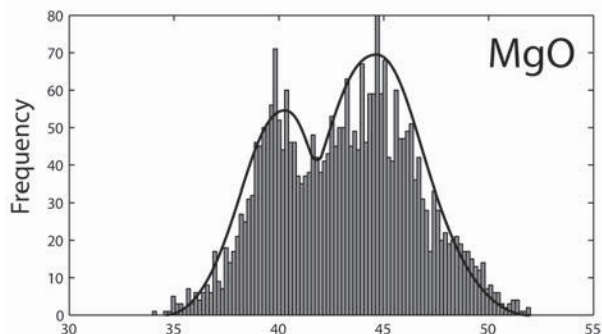
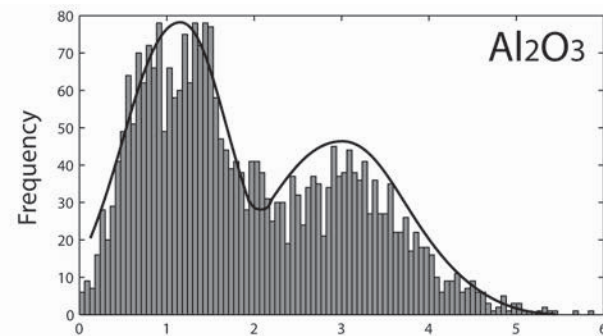
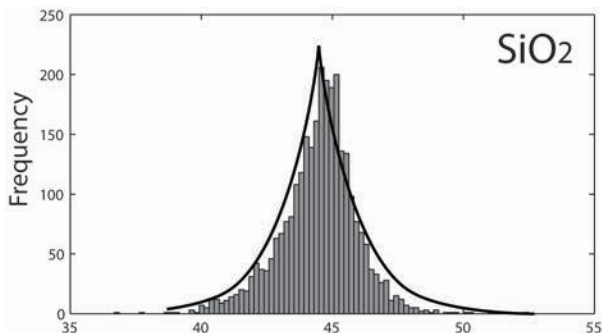


$$\rho(\mathbf{d}, \mathbf{m}) = \text{const.} \exp \left\{ -\frac{1}{2} (\mathbf{d} - \mathbf{g}(\mathbf{m}))^T \mathbf{C}_t^{-1} (\mathbf{d} - \mathbf{g}(\mathbf{m})) \right\}$$

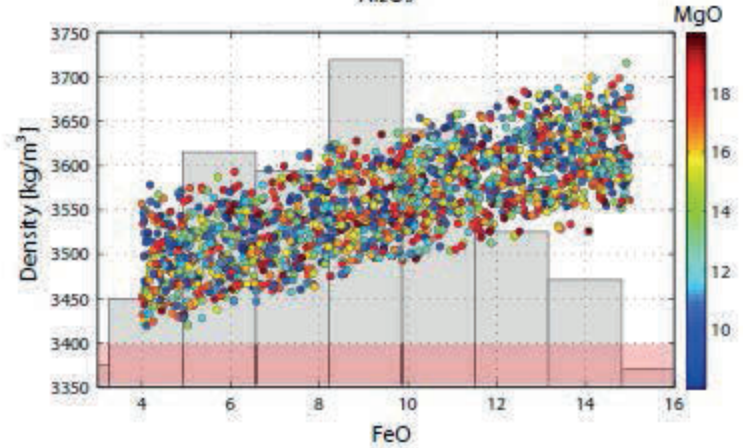
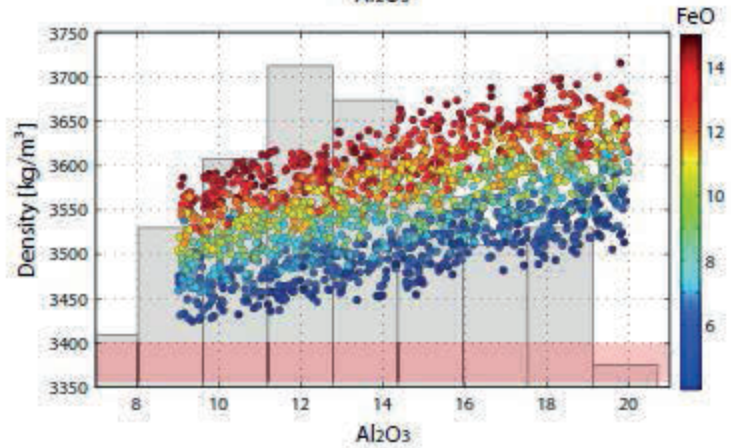
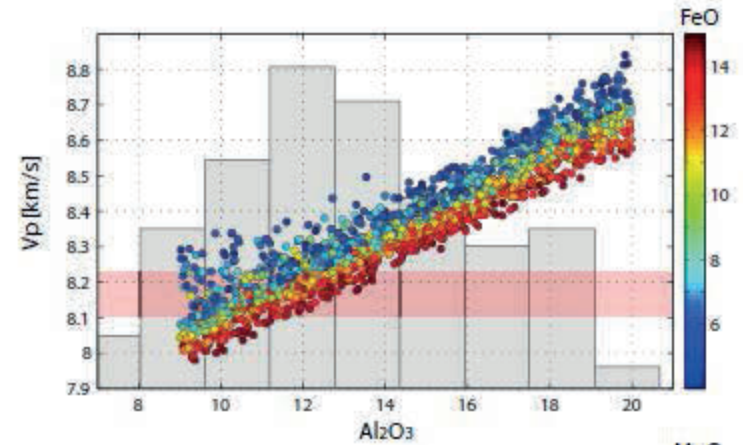
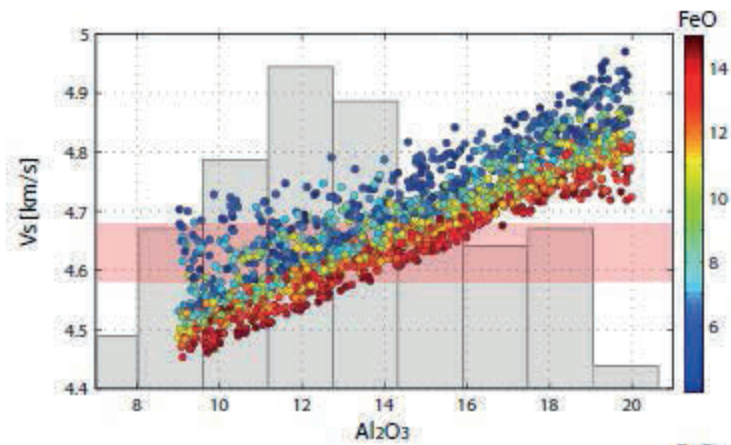


2) How do we do it?

Be careful!! the distribution can be biased!



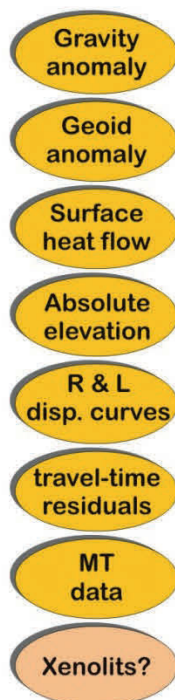
The two peaks in our distributions could be an artifact of the sampling



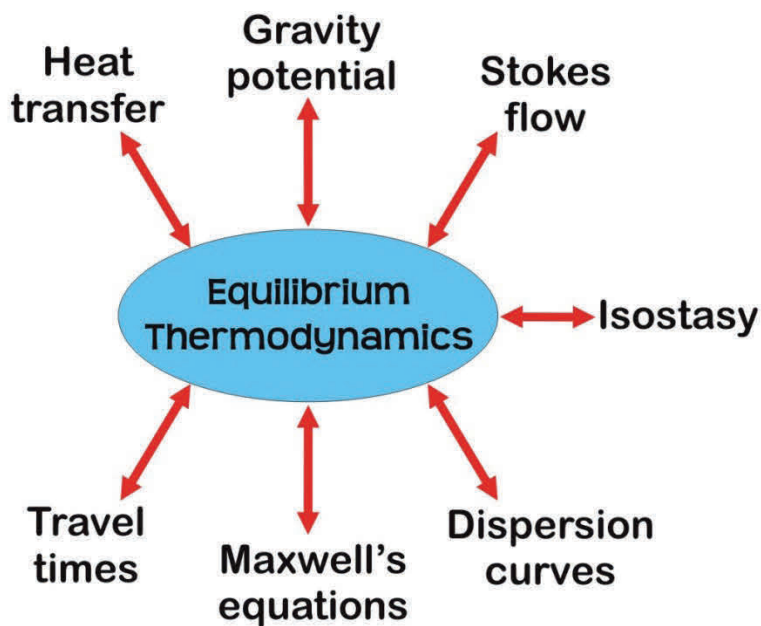
3) Our approach ...so far

LitMod_4INV

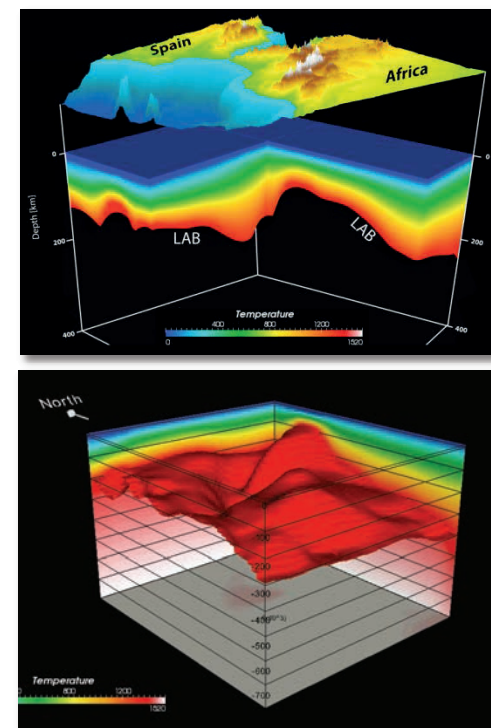
DATA



THEORY



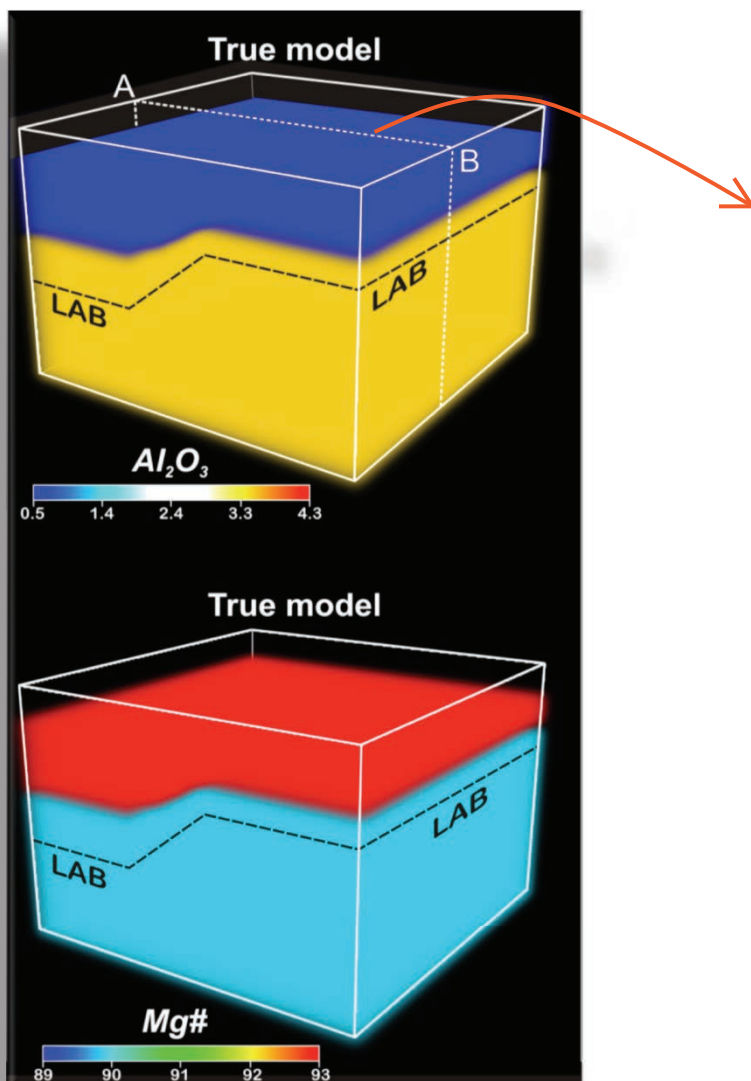
MODELS



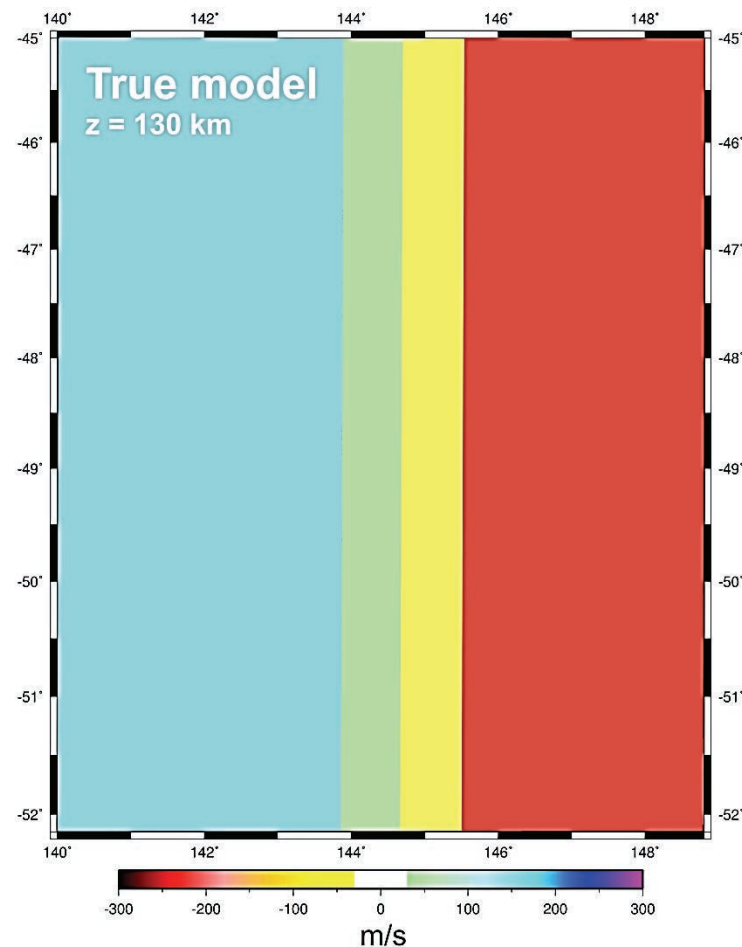
Probabilistic Inversion
(Bayesian framework)

3) Our approach ...so far

Let's look at a simple example...

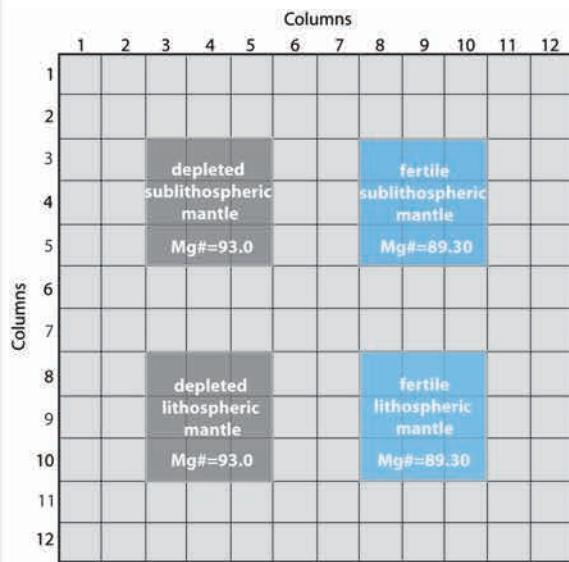


Vp anomalies relative to horizontal average

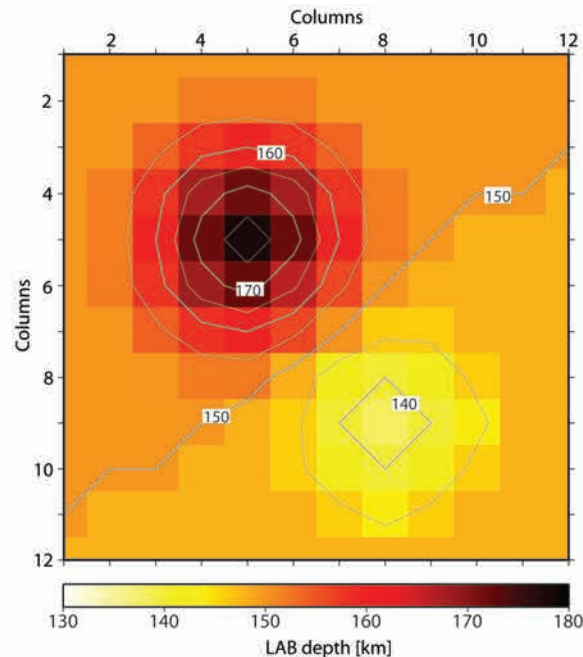


4) A synthetic example

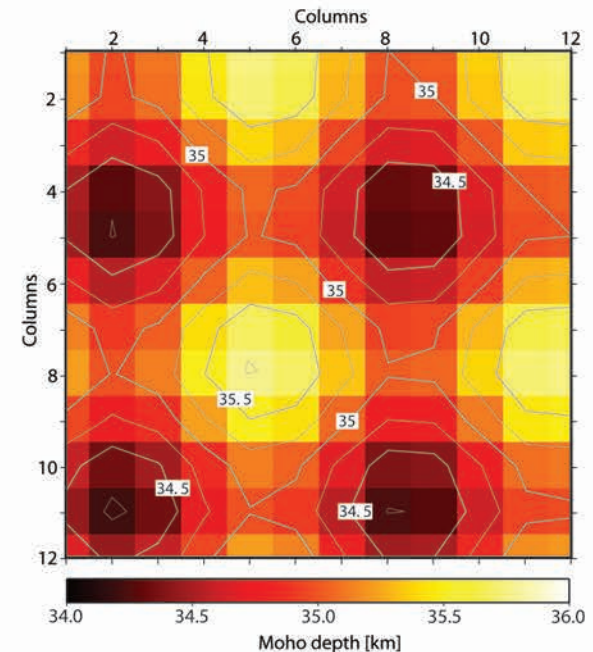
We created a 3D model with *LitMod3D* and used its predictions + noise as input data for the inversion



Domains



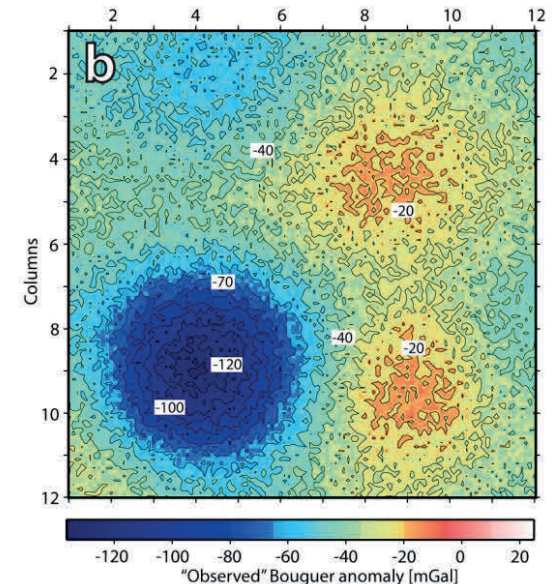
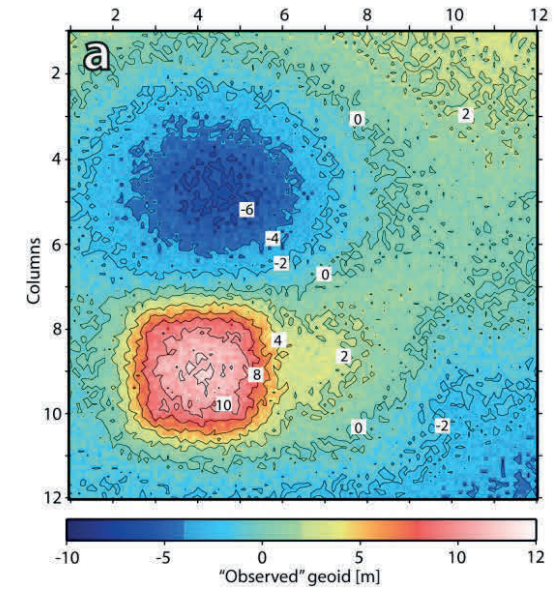
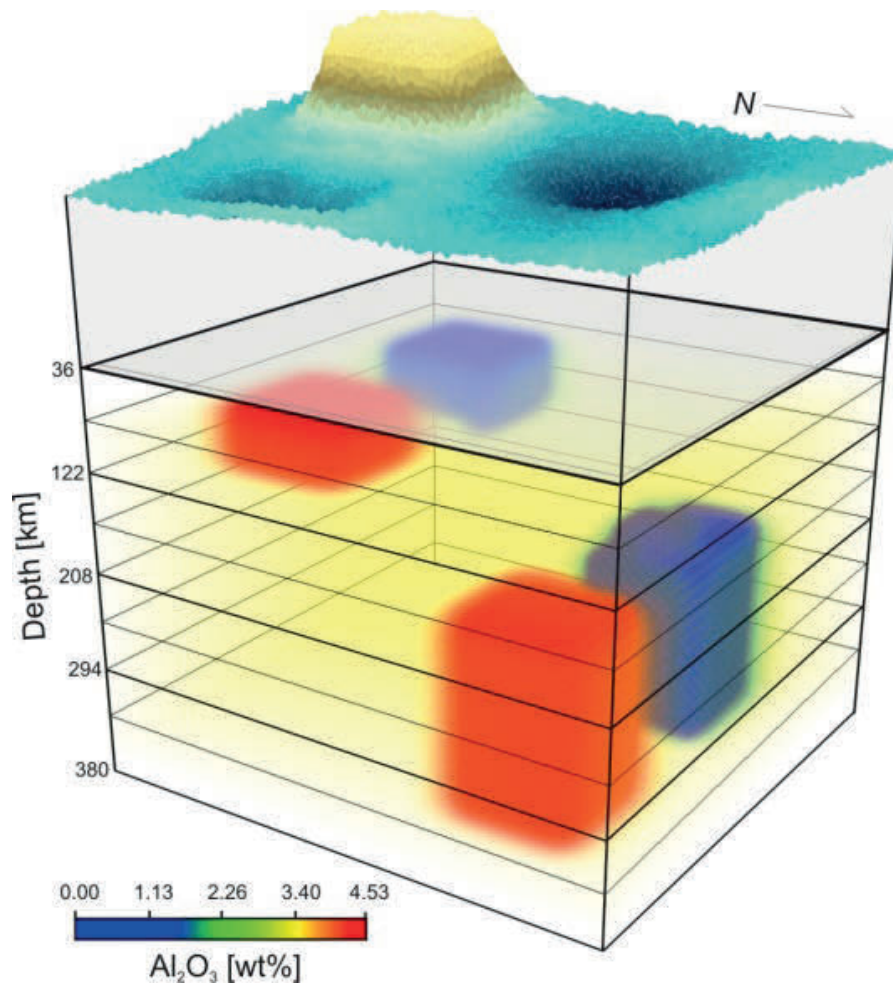
LAB



Moho

4) A synthetic example

Examples of inputs (predictions + noise) used in the inversion



4) A synthetic example

Results for LAB geometry

Mean models of 8 random ensembles with 500 samples each taken from the total posterior

*Note the
“persistent”
features*

